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Links between seawater flooding, soil ammonia oxidiser communities and their response to changes in salinity

Journal:	<i>FEMS Microbiology Ecology</i>
Manuscript ID	FEMSEC-17-04-0166.R2
Manuscript Type:	Research article
Date Submitted by the Author:	18-Sep-2017
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Keywords:	ammonia-oxidising archaea, ammonia-oxidising bacteria, nitrification, flooding, salinity, <i>Nitrosococcus</i>

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104x68mm (300 x 300 DPI)

Review

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3 **1 Links between seawater flooding, soil ammonia oxidiser communities and**
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5 **2 their response to changes in salinity**
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16 **Keywords:** ammonia-oxidising archaea; ammonia-oxidising bacteria; nitrification; flooding;
17 salinity; *Nitrosococcus*
18

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23

24 **Abstract**

25 Coastal areas worldwide are challenged by climate change-associated increases in sea level
26 and storm surge quantities that potentially lead to more frequent flooding of soil ecosystems.
27 Currently, little is known of the effects of inundation events on microorganisms controlling
28 nitrification in these ecosystems. The goal of this study was to investigate the impact of
29 seawater flooding on the abundance, community composition and salinity tolerance of soil
30 ammonia oxidisers. Topsoil was sampled from three islands flooded at different frequencies
31 by the Wadden Sea. Archaeal ammonia oxidiser *amoA* genes were more abundant than
32 their betaproteobacterial counterparts and the distribution of archaeal and bacterial ammonia
33 oxidiser *amoA* and 16S rRNA gene sequences significantly differed between the islands.
34 The findings indicate selection of ammonia oxidiser phylotypes with greater tolerance to high
35 salinity and slightly alkaline pH (e.g., *Nitrosopumilus* representatives) in frequently flooded
36 soils. A cluster phylogenetically related to gammaproteobacterial ammonia oxidisers was
37 detected in all samples analysed in this survey. Nevertheless, no gammaproteobacterial
38 *amoA* genes could be amplified via PCR and only betaproteobacterial ammonia oxidisers
39 were detected in enrichment cultures. A slurry-based experiment demonstrated tolerance of
40 both bacterial and archaeal ammonia oxidisers to a wide range of salinities (e.g., Wadden
41 Sea water salinity) in soil naturally exposed to seawater at a high frequency.

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48 Introduction

49 Flooding of the Earth's low-lying coastal land is expected to increase as a result of recent
50 and predicted future increases in sea levels (Nicholls and Cazenave 2010; Haigh *et al.* 2014;
51 Wong *et al.*, 2014) and estimated higher number and intensification of storm surge events
52 (Dasgupta *et al.* 2009; Mousavi *et al.* 2011; Grinsted *et al.* 2013). Soil ecosystems that are
53 directly exposed to seawater during inundation are particularly susceptible to the effects of
54 flooding, such as osmotic stress. It has been reported that tsunami-induced seawater
55 inundation significantly alters soil bacterial community structure (Asano *et al.* 2013). In
56 addition, soil salinity, which can be affected by flooding events, was found to be strongly
57 related to bacterial community structure in tidal wetlands (Morrissey *et al.* 2014). Currently,
58 the impacts of temporal seawater inundation on specific functional microbial groups,
59 essential for soil fertility and globally important biogeochemical cycles are poorly understood.

60 The initial step of nitrification, a key process within the nitrogen cycle, is mediated by
61 ammonia oxidisers (AO) (Prosser and Nicol 2008), which comprise ammonia oxidising
62 archaea (AOA) and bacteria (AOB). AOA outnumber their bacterial counterparts in various
63 marine environments such as North Sea or Atlantic Ocean water (Wuchter *et al.* 2006).
64 Furthermore, AOA and AOB cohabit most terrestrial ecosystems (Prosser 2011), but their
65 relative abundances in soil show significant variations (Prosser and Nicol 2012). Previous
66 studies indicate that soil pH represents a major driver of these variations, with AOA typically
67 dominating acidic niches (Nicol *et al.* 2008; Gubry-Rangin *et al.* 2010; Yao *et al.* 2013).
68 There is also limited evidence for greater sensitivity of AOA toward high ammonia
69 concentration and their preference of soils with low ammonia supply (Prosser and Nicol
70 2012). Oxygen availability represents another parameter affecting AO communities and few
71 studies indicate that AOA have a higher affinity for oxygen than AOB (Chen *et al.* 2008; Jung
72 *et al.* 2011; Pitcher *et al.* 2011; Kim *et al.* 2012). In wet tropical soil, which experiences
73 fluctuating oxygen availability, AOA tolerant of extended periods of anoxia were detected
74 (Pett-Ridge *et al.* 2013). Moreover, relationships between salinity and sediment AOA or AOB

1
2
3 75 distribution were reported (Coci *et al.* 2005; Sahan and Muyzer 2008; Bernhard *et al.* 2010;
4
5 76 Zheng *et al.* 2014; Zhang *et al.* 2015). For instance, it has been shown that salinity is a
6
7 77 steering factor in selecting *Nitrosomonas* and *Nitrosopumilus* representatives (Coci *et al.*
8
9 78 2005; Zhang *et al.* 2015).

10
11
12 79 A study based on high-throughput sequencing of *amoA* genes (encoding subunit A of
13
14 80 ammonia monooxygenase) indicates that AOA are dominated by the *Nitrososphaera* cluster
15
16 81 in numerous soil ecosystems, with the exception of acidic niches (Pester *et al.* 2012). In
17
18 82 contrast, members of the *Nitrosopumilus* cluster (also called thaumarchaeal group I.1a or
19
20 83 marine group) represent the dominant AOA in oligotrophic seawater, which is probably due
21
22 84 to high affinity for ammonia (Martens-Habben *et al.* 2009). AOB comprise the
23
24 85 betaproteobacterial genera *Nitrosomonas* and *Nitrosospira*, collectively forming the family
25
26 86 *Nitrosomonadaceae*, as well as the gammaproteobacterial genera *Nitrosococcus* and
27
28 87 *Nitrosoglobus*. Different *Nitrosomonas* lineages, such as the *Nitrosomonas marina* lineage
29
30 88 and the *Nitrosomonas europaea/Nitrosomonas mobilis* lineage, are predominantly detected
31
32 89 in marine or non-marine environments (Prosser 2014). *Nitrosococcus* has been poorly
33
34 90 studied with molecular techniques but cultured representatives were exclusively derived from
35
36 91 salty environments (Campbell *et al.* 2011; Wang *et al.* 2017). Furthermore, the
37
38 92 gammaproteobacterial AOB *Nitrosoglobus terrae* TAO100 was recently isolated from an
39
40 93 acidic agricultural soil (Hayatsu *et al.* 2017).

41
42
43 94 Both, AOA and AOB, have been detected in coastal soil and sediment (Moin *et al.* 2009;
44
45 95 Marton *et al.* 2015; Bernhard *et al.* 2016). In New England salt marsh sites, up to
46
47 96 approximately 10^9 AOA *amoA* genes g^{-1} dry sediment were reported (Moin *et al.* 2009) and
48
49 97 higher abundances of AOA than AOB *amoA* genes were found in soil and sediment samples
50
51 98 from salt marshes (Moin *et al.* 2009; Marton *et al.* 2015). Nevertheless, information about
52
53 99 effects of different seawater flooding frequencies on ammonia oxidisers inhabiting coastal
54
55 100 soil is lacking. As representatives of AOA and AOB show salinity preferences (Koops *et al.*
56
57 101 1991; Campbell *et al.* 2011; Qin *et al.* 2014), it can be assumed that both groups are
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1
2
3 102 affected by inundation events. The German Halligen can be considered a model system for
4
5 103 extensive analysis of flooding effects. These islands are surrounded by the Wadden Sea and
6
7 104 are inundated at different frequencies during the year. Approximately 100 years ago, dykes
8
9 105 were established for the Halligen Hooge and Langeness (Schindler *et al.* 2014a). Due to the
10
11 106 dykes and closing tidal gates with rising tides, Hooge and Langeness need much higher
12
13 107 gauge levels to be flooded than Hallig Nordstrandischmoor. Therefore, Nordstrandischmoor
14
15 108 shows higher seawater inundation frequencies than the Halligen Hooge and Langeness.
16

17
18 109 In this study, the impacts of seawater flooding on abundance, diversity and community
19
20 110 composition of soil AO were investigated. Flooding significantly affects physicochemical
21
22 111 parameters such as salinity or oxygen level in soil. Here, we focused on salinity, which can
23
24 112 show strong fluctuations in temporarily flooded soil over time. In order to explore the
25
26 113 response of ammonia oxidisers to changes in sea salt concentration, a soil slurry experiment
27
28 114 was performed. All analyses were carried out using topsoil samples derived from Hooge,
29
30 115 Langeness and Nordstrandischmoor. The study was based on the following hypotheses: (1)
31
32 116 flooding of soil ecosystems evokes changes in AOA and AOB abundance due to changes in
33
34 117 salinity and/or invasion of AO in surrounding seawater, (2) representatives of AO inhabiting
35
36 118 soil frequently inundated with seawater exhibit halotolerance and tolerance toward changing
37
38 119 salt concentrations, (3) the abundance of thaumarchaeal group I.1a (marine group) is
39
40 120 positively correlated with seawater flooding frequency and (4) within AOB, phylogenetic
41
42 121 lineages typically encountered in marine habitats show higher abundance in frequently
43
44 122 flooded soil than in soil rarely inundated with seawater.
45

46 47 123 **Materials and Methods**

48 49 50 124 **Study sites and soil sampling**

51
52
53 125 The German Halligen, ten small islands with a total area of 23 km², are surrounded by the
54
55 126 North Frisian Wadden Sea. They represent relics of former mainland or larger islands. At the
56
57 127 beginning of the 20th century, dykes were established to strengthen the marshland of Hooge
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1
2
3 128 and Langeness against storm surges. The dykes show an average height of 1.5 m (Hooge)
4
5 129 and 1.0 m (Langeness) above the present 10-year (2001-2010) mean high water (Schindler
6
7 130 *et al.* 2014b). Inundation of Hooge and Langeness marshland only occurs during heavy,
8
9 131 westerly storm conditions (Schindler *et al.* 2014b). In contrast, the marshland of
10
11 132 Nordstrandischmoor is not enclosed by a dyke and therefore exhibits a high inundation
12
13 133 frequency, with an average of 15 events year⁻¹ between 2001 and 2010 (Schindler *et al.*
14
15 134 2014b), while the ten-year average (2001-2010) for Hooge is two events year⁻¹ (Schindler *et*
16
17 135 *al.* 2014b). The flooding frequency of Langeness (10 events year⁻¹) is intermediate.

18
19
20 136 Topsoil was sampled from Hooge (Ho), Langeness (La) and Nordstrandischmoor (No)
21
22 137 (Figure 1), with coordinates of sampling sites (Ho1, Ho2, Ho3, La1, La2, La3, La4, No1, No2,
23
24 138 and No3) provided in Table S1. Triplicate samples at each sampling site (designated, e.g.,
25
26 139 Ho1a, Ho1b, and Ho1c) were removed at 1-m distances from each other. The upper 10 cm
27
28 140 of the mineral soil was separated from adjacent soil material, homogenised and sieved to
29
30 141 <3.5 mm. Soil samples for slurry-based analyses were stored at 4°C and all other samples
31
32 142 were kept frozen (-80°C).

33 34 35 143 **Basic soil analyses and DNA extraction**

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37
38 144 Subsamples (2 g) of field-moist soils were used to determine gravimetric water content by
39
40 145 drying to constant weight at 105°C. Another subsample (5 g) was extracted in 25 ml 1 M KCl
41
42 146 for analysis of ammonium (NH₄⁺), nitrite (NO₂⁻) and nitrate (NO₃⁻) using a flow injection
43
44 147 analyser "Quikchem QC85S5" (Lachat Instruments, Loveland, CO, USA). Air-dried soil was
45
46 148 used to determine soil pH, electrical conductivity and C and N concentrations. Soil pH was
47
48 149 measured in a 1:2.5 suspension of soil in distilled water. Electrical conductivity was
49
50 150 measured in the same extract using a "Tetra Con 325" electrode (WTW, Weilheim,
51
52 151 Germany). Ground subsamples were analysed for total carbon (TC) and nitrogen (TN)
53
54 152 concentrations by dry combustion with a CN analyser "Vario Max" (Elementar
55
56 153 Analysensysteme GmbH, Hanau, Germany). Inorganic carbon (IC) was determined after
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3 154 combustion of organic carbon in a muffle furnace (450°C for 16 h) and soil organic carbon
4
5 155 (SOC) was calculated as the difference between TC and IC.
6
7

8 156 Genomic DNA was extracted from 0.5 g soil using the PowerSoil DNA isolation kit (MoBio
9
10 157 Laboratories, Carlsbad, CA, USA) according to the manufacturer's instruction. The
11
12 158 concentration of extracted DNA was estimated using a NanoDrop 1000 spectrophotometer
13
14 159 (Thermo Fisher Scientific, Loughborough, UK).
15

160 **Soil slurry incubations**

161 Slurries containing soil derived from Hooge, Langeness, or Nordstrandischmoor and artificial
162 seawater were prepared in 500-ml Duran bottles and consisted of 20 g soil (moisture content
163 adjusted to 45% with sterile distilled water) in 100 ml artificial seawater. A wide range of
164 salinity conditions were covered using artificial seawater (Subow 1931) with different salt
165 concentrations (5 ‰, 10 ‰, 20 ‰, or 30 ‰). Ammonia was not added to prevent inhibition of
166 ammonia-sensitive ammonia oxidizers by unnaturally high concentrations of ammonia.
167 Triplicate slurries were incubated in the dark for 30 days at 26°C. Every third day, slurries
168 were shaken and seals were removed for 5 min. Samples (8 ml) were collected following
169 shaking, using a pipette with a 3-mm diameter tip, directly after establishing slurries and after
170 incubation for 15 and 30 days. Samples were centrifuged in Falcon tubes at 18,500 × g and
171 4°C for 5 min and pellets were stored frozen until DNA extraction and supernatants were
172 used for analysis of NH_4^+ , NO_2^- and NO_3^- . Duran bottles containing soil slurries were
173 weighed several times during incubation to calculate evaporation, which was found to be
174 negligible.
175

175 **Quantitative real-time PCR analysis of *amoA* genes**

176 Quantitative real-time PCR (qPCR) was carried out using a BioRad MyIQ Single-Color Real-
177 Time PCR Detection System (BioRad, Hertfordshire, UK) and SybrGreen as fluorescent dye.
178 Primer sets CrenamoA23f/CrenamoA616r (Tourna *et al.* 2008) and *amoA*-1F/*amoA*-2R
179 (Rotthauwe *et al.* 1997) were used for AOA and betaproteobacterial AOB *amoA*
180

1
2
3 180 quantification, respectively. The PCR mixture (final volume, 25 μ L per reaction) contained
4
5 181 25 μ g bovine serum albumin (BSA), 400 nM of each AOA or AOB *amoA*-specific primer,
6
7 182 12.5 μ L of QuantiFast™ qPCR master mix (Qiagen, Crawley, UK), and 25 ng of DNA. The
8
9 183 following cycling conditions were used: 15 min at 95°C, 40 cycles of 15 s at 94°C, 1 min 30 s
10
11 184 for AOA *amoA* and 1 min for AOB *amoA* at 60°C and plate fluorescence measurement after
12
13 185 8 s at 80°C. Melting curve analysis and agarose gel electrophoresis were performed to
14
15 186 assess specificity of amplification during PCR. Amplification efficiency ranged from 87.6% to
16
17 187 102.5% and from 80.3% to 89.7% for AOA and betaproteobacterial AOB *amoA*, respectively,
18
19 188 with r^2 values >0.98.

189 **Pyrosequencing of bar-coded *amoA* and 16S rRNA gene amplicons**

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24
25 190 AOA *amoA* genes and the V3-V5 region of bacterial 16S rRNA genes were amplified by
26
27 191 PCR using DNA extracts from the collected soil samples as template (samples derived from
28
29 192 soil slurry incubations were not considered). Primer CrenamoA23f (Tourna *et al.* 2008),
30
31 193 containing a sample-specific MID (extended multiplex identifier) and Roche 454
32
33 194 pyrosequencing adaptor A, as well as primer CrenamoA616r (Tourna *et al.* 2008), containing
34
35 195 Roche 454 pyrosequencing adaptor B, were used for amplification of AOA *amoA*. The
36
37 196 forward and reverse primers V3for 5'-CCATCTCATCCCTGCGTGTCTCCGACTCAG-MID-
38
39 197 TACGGRAGGCAGCAG-3' and V5rev 5'-
40
41 198 CCTATCCCCTGTGTGCCTTGGCAGTCTCAGCCGTCAATTCMTTGTGAGT-3' (Roche 454
42
43 199 pyrosequencing adaptors are underlined) (Liu *et al.* 2007; Wang and Qian 2009; Nacke *et al.*
44
45 200 2016) were used to amplify the V3-V5 region of bacterial 16S rRNA genes. The PCR
46
47 201 reaction mixture (50 μ L) contained 10 μ L 5-fold reaction buffer (Phusion GC buffer; Thermo
48
49 202 Fisher Scientific), 200 μ M of each of the four deoxynucleoside triphosphates, 5% dimethyl
50
51 203 sulfoxide (DMSO), 0.2 μ M of each of the primers, 1 U Phusion high-fidelity DNA polymerase
52
53 204 (Thermo Fisher Scientific) and 50 ng of DNA. The following PCR cycling conditions were
54
55 205 used for amplification of AOA *amoA*: initial denaturation at 98°C for 30 s, 30 cycles of
56
57 206 denaturation at 98°C for 10 s, annealing at 55°C for 30 s and extension at 72°C for 30 s,
58
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1
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3 207 followed by a final extension step at 72°C for 5 min. PCR reactions for amplification of
4
5 208 bacterial 16S rRNA genes were initiated at 98°C (30 s), followed by 25 cycles of 98°C (10 s),
6
7 209 65°C (30 s) and 72°C (30 s), and ended with incubation at 72°C for 5 min. All samples were
8
9 210 amplified in triplicate, purified using the peqGold gel extraction kit (Peqlab Biotechnologie
10
11 211 GmbH, Erlangen, Germany) and pooled in equal amounts. Sequences of *amoA* and 16S
12
13 212 rRNA gene amplicons were determined using a Roche GS-FLX 454 pyrosequencer and
14
15 213 titanium chemistry as recommended by the manufacturer.

18 214 **Pyrosequencing data processing and analysis**

20
21 215 AOA *amoA* sequences exhibiting an average quality value of <30, primer mismatches or
22
23 216 long homopolymers (>8 bp) were removed using QIIME (version 1.7) (Caporaso *et al.* 2010).
24
25 217 In addition, reads were truncated at the first ambiguous base encountered. After truncation
26
27 218 of remaining primer sequences, noise resulting from amplicon pyrosequencing was removed
28
29 219 using the Acacia error-correction tool (version 1.5) (Bragg *et al.* 2012). Chimeras were
30
31 220 removed using Uchime (version 7.0) (Edgar *et al.* 2011) and using 735 *amoA* sequences
32
33 221 comprising the currently known five major clusters of archaeal *amoA* as a reference
34
35 222 database, obtained from the ARB (Ludwig *et al.* 2004) database created by Pester *et al.*
36
37 223 (2012). Putative chimeras identified by Uchime were compared with the NCBI non-redundant
38
39 224 database using BLASTn. Clear breaks in the local alignments affiliating queried *amoA*
40
41 225 amplicons to two very distinct sequence entries in the NCBI database confirmed Uchime-
42
43 226 based chimera detection. Sequences classified as chimeras by Uchime and manual
44
45 227 inspection of local alignments were discarded completely. Subsequently, all sequences
46
47 228 shorter than 592 bp were removed. Determination of operational taxonomic units (OTUs) at
48
49 229 85% genetic similarity was performed using Uclust (Edgar 2010). Singleton OTUs containing
50
51 230 only one *amoA* sequence were removed using QIIME. A representative sequence for each
52
53 231 OTU was added to the consensus tree for archaeal *amoA* phylogeny constructed by Pester
54
55 232 *et al.* (2012) using the parsonomy tool in ARB (Ludwig *et al.* 2004). Furthermore, *amoA*
56
57 233 sequences of *Nitrosocosmicus franklandus* (Lehtovirta-Morley *et al.* 2016), *Nitrosoarchaeum*

1
2
3 234 *koreensis* MY1 (Jung *et al.* 2011), and *Nitrosotalea devanaterrea* Nd1 (Lehtovirta-Morley *et al.*
4
5 235 2011) were added to this tree. OTU-based analyses were performed at the same level of
6
7 236 surveying effort (1,200 randomly selected sequences per sample). Calculation of the Chao1
8
9 237 index (Chao and Bunge 2002) and the Shannon index (Shannon 1948) was conducted using
10
11 238 QIIME.

12
13
14 239 Bacterial 16S rRNA gene sequences were pre-processed as described by Broszat *et al.*
15
16 240 (2014), except that sequences shorter than 300 bp were removed using QIIME. OTUs were
17
18 241 determined at a genetic distance of 3% using Uclust (Edgar 2010), implemented in QIIME
19
20 242 (Caporaso *et al.* 2010), and were classified by comparison of partial 16S rRNA gene
21
22 243 sequences using the SILVA SSU database (Pruesse *et al.* 2007) release 119. OTUs
23
24 244 classified as chloroplast or mitochondrion as well as unclassified OTUs were removed from
25
26 245 16S rRNA gene sequence datasets and rarefaction curves were generated using QIIME. A
27
28 246 Neighbor Joining phylogenetic tree was constructed in MEGA (version 7.0) (Kumar *et al.*
29
30 247 2016) using the Maximum Composite Likelihood model based on a ClustalW alignment of
31
32 248 sequences (truncated to a length of 457 bp) affiliated to selected OTUs, uncultured
33
34 249 *Gammaproteobacteria* closely related to these OTUs, *Chromatiaceae*, and *Acidithiobacillus*.
35
36 250 A total number of 500 bootstrap resamplings were carried out to test the tree topology.

37 38 39 251 **Statistical analysis**

40
41
42 252 Data were tested for normal distribution before performance of ANOVA. One-way ANOVA
43
44 253 and Tukey pair-wise comparisons were conducted to determine differences in basic soil
45
46 254 properties, *amoA* abundance and Shannon diversity between Hooge, Langeness, and
47
48 255 Nordstrandischmoor. The effect of salinity on *amoA* abundance was determined by ANOVA
49
50 256 of data from triplicate slurries followed by Tukey *post hoc* tests. The PAST statistical
51
52 257 package (Hammer *et al.* 2001) was used to perform Mann-Whitney U tests to identify taxa
53
54 258 showing significant differences in relative abundance between Hallig soils. Non-metric
55
56 259 multidimensional scaling was performed based on weighted Unifrac (Lozupone *et al.* 2011)

1
2
3 260 distance matrices (genetic distance: 3%) using the vegan package (Oksanen *et al.* 2016) in
4
5 261 R (R Development Core Team, 2015).
6
7

8 262 **Enrichment of AO and analysis of cultivates**

9
10 263 In order to enrich gammaproteobacterial AOB, potentially inhabiting temporarily flooded soil,
11
12 264 sea salts-containing medium has been inoculated with sample material derived from the
13
14 265 Halligen. Incubation was carried out in 100-mL Duran bottles containing 50 mL enrichment
15
16 266 medium without shaking at 28°C in the dark. Per bottle, 1 g of topsoil was used as inoculum.
17
18 267 After autoclaving of 30% or 50% artificial seawater described by Stephen *et al.* (1996), the
19
20 268 following components were added aseptically (per liter) to obtain cultivation medium: 2 ml
21
22 269 1 M NaHCO₃, 1 ml 7.5 mM EDTA-Na₂, 1 ml non-chelated trace element mixture (Widdel and
23
24 270 Bak, 1992), 1 ml selenite-tungstate solution (Widdel and Bak, 1992), 1 ml vitamin solution
25
26 271 (Balch *et al.* 1979), and 0.2 g KH₂PO₄. Furthermore, 500 µM or 12.5 mM NH₄Cl was added.
27
28 272 The pH of the medium was adjusted to 7.8 by the addition of 5% (wt/vol) Na₂CO₃ before
29
30 273 filter-sterilization through a 0.22 µm pore size bottle-top filter. Subculturing was carried out
31
32 274 by transferring 5% culture into fresh enrichment medium. The pH was monitored and if
33
34 275 necessary adjusted. Growth of ammonia oxidizers was assessed by analysis of inorganic
35
36 276 nitrogen. Colorimetric ammonia and nitrite measurement was performed as described by
37
38 277 Lehtovirta-Morley *et al.* (2014). Enrichment cultures were continuously subcultured in the
39
40 278 laboratory for approximately two years.
41
42
43

44 279 To analyse enriched bacteria, DNA was extracted and the V3-V4 region of 16S rRNA genes
45
46 280 amplified by PCR. Primers S-D-Bact-0341-b-S-17 and S-D-Bact-0785-a-A-21 (Klindworth *et*
47
48 281 *al.* 2013) containing modifications for Illumina MiSeq sequencing were employed. The PCR
49
50 282 reaction mixture was prepared as described above for amplification of the V3-V5 region of
51
52 283 bacterial 16S rRNA genes but it contained 3% instead of 5% DMSO. PCR reactions were
53
54 284 initiated at 98°C (1 min), followed by 29 cycles of 98°C (45 s), 60°C (45 s) and 72°C (30 s),
55
56 285 and ended with incubation at 72°C for 5 min. PCR products were purified using the
57
58
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1
2
3 286 NucleoMag 96 PCR kit (Macherey and Nagel, Düren, Germany) and sequences determined
4
5 287 using an Illumina MiSeq instrument. Demultiplexing and clipping of sequence adapters from
6
7 288 raw sequences was performed by employing CASAVA data analysis software (Illumina).
8
9 289 Paired-end sequences were merged using PEAR v0.9.10 (Zhang *et al.* 2014) with default
10
11 290 parameters. Subsequently, sequences with an average quality score lower than 20 and
12
13 291 containing unresolved bases were removed with the `split_libraries_fastq.py` script from
14
15 292 QIIME (Caporaso *et al.* 2010). Remaining primer sequences were truncated using the
16
17 293 `cutadapt` program (Martin, 2011) and we employed Usearch (Edgar, 2010) to sort out
18
19 294 sequences shorter than 400 bp. Chimeric sequences were removed using Uchime (Edgar *et*
20
21 295 *al.* 2011) in reference mode against the Greengenes Gold data set
22
23 296 `gold_strains_gg16S_aligned.fasta` (DeSantis *et al.* 2006). OTUs were determined at a
24
25 297 genetic distance of 3% using Uclust (Edgar 2010) and classified by comparison of partial
26
27 298 16S rRNA gene sequences with the SILVA SSU database (Pruesse *et al.* 2007).

299 **Sequence data deposition**

300 Sequence data were deposited in the Sequence Read Archive (SRA) of the National Center
31
32
33 301 for Biotechnology Information (NCBI) under the accession numbers SRP083741 and
34
35 302 SRP114677.

303 **Results**

304 **Soil characteristics and ammonia oxidiser abundance**

305 Salinity, measured as electrical conductivity, and pH were significantly higher in soils from
36
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42
43 306 Nordstrandischmoor than in soils from Langeness or Hooge (Table 1). This is likely due to
44
45
46 307 the high frequency of flooding of Nordstrandischmoor soil with seawater with a pH of
47
48
49 308 approximately 8.1, increasing the soil pH to 7.1 - 7.9. Total N, organic C and total C did not
50
51
52 309 differ significantly between the three soils, but ammonium and nitrate concentrations were
53
54
55 310 significantly different (Table 1). Furthermore, ammonium concentrations were higher than
56
57
58 311 nitrate concentrations at each site (Table 1).

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2
3 312 AOA and betaproteobacterial AOB were detected in all Hallig soil samples by qPCR
4
5 313 amplification of *amoA* genes (Figure 2), with AOA and AOB *amoA* abundances, respectively,
6
7 314 ranging from 2.85×10^6 to 1.37×10^8 genes g^{-1} dry soil and 1.86×10^5 to 6.11×10^7 genes
8
9 315 g^{-1} dry soil. AOA *amoA* genes were significantly more abundant than betaproteobacterial
10
11 316 AOB *amoA* genes ($P < 0.05$) but there was no obvious difference between the three Halligen
12
13 317 (Figure 2), although abundances of both genes were higher in the Nordstrandischmoor soil
14
15 318 than in soil from Hooge ($P < 0.001$). No significant relationships between nitrate data and
16
17 319 AOA or AOB *amoA* gene abundance were detected.

20 320 **Response of ammonia oxidisers to salt concentration**

21
22
23 321 A soil slurry experiment was performed to assess whether soil AO inhabiting Hooge,
24
25 322 Langeness and Nordstrandischmoor, flooded by seawater at different frequencies, respond
26
27 323 differentially to salinity. AOA *amoA* abundance in Nordstrandischmoor soil slurries
28
29 324 significantly increased during incubation at 5‰ and 10‰ salinity ($P < 0.05$, Figure 3a). In
30
31 325 addition, average AOA *amoA* abundance increased at the two highest salinities tested
32
33 326 (Figure 3a). In contrast, AOA *amoA* abundance in slurries containing Hooge or Langeness
34
35 327 soil significantly decreased after incubation with 30‰ salinity for 30 days ($P < 0.05$, Figure
36
37 328 3a). There were no significant effects of time or salinity on betaproteobacterial AOB *amoA*
38
39 329 abundance in Nordstrandischmoor soil slurries, but there was evidence of decreases in
40
41 330 abundance in Langeness soil slurries, which were significant at 5‰ and 30‰ salinity (Figure
42
43 331 3b). In Hooge soil slurries, the only significant change was an increase in abundance at 20‰
44
45 332 after incubation for 30 days.

46
47
48 333 Changes in inorganic nitrogen in soil slurries are presented in Figure 4. Ammonium
49
50 334 concentration in Hooge and Langeness soil slurries was significantly greater after incubation
51
52 335 for 15 days than at 0 days. This increase is potentially due to mineralisation of existing
53
54 336 organic nitrogen. The smallest increase was detected in Hooge and Langeness soil slurries
55
56 337 containing artificial seawater with a salt concentration of 20 ‰, indicating again a preference
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1
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3 338 for this salinity. Furthermore, combined nitrite plus nitrate concentrations significantly
4
5 339 increased between day 15 and day 30 of incubation in Hooge slurries with an artificial
6
7 340 seawater salinity of 20 ‰ (Figure 4). In contrast, ammonium concentration in
8
9 341 Nordstrandischmoor slurries was significantly lower after incubation for 15 days than at 0
10
11 342 days at artificial seawater salinities of 10 ‰, 20 ‰ and 30 ‰.

14 343 **Flooding effects on ammonia oxidiser diversity and community composition**

16
17 344 Pyrosequencing of AOA *amoA* gene amplicons yielded 76,427 raw sequences, of which
18
19 345 54,122 sequences were used following quality control. These sequences contained eight
20
21 346 OTUs, defined at 85% AOA *amoA* gene sequence identity, which has been suggested as a
22
23 347 'species' threshold for AOA based on phylogenetic analysis of a comprehensive set of
24
25 348 publicly available *amoA* sequences (Pester *et al.* 2012). Comparison of rarefaction analyses
26
27 349 with the number of OTUs calculated by the Chao1 richness estimator indicate that estimated
28
29 350 richness was covered by the sequencing effort in 27 of the 30 analysed samples, with
30
31 351 approximately 80% coverage in the remaining three samples, and individual samples
32
33 352 contained 3 - 5 OTUs (Table S2). Shannon diversities for OTUs in the soils from the different
34
35 353 Halligen were not significantly different ($P=0.93$, Table S2).

37
38 354 The dominant AOA OTUs were affiliated to *Nitrosopumilus* subcluster 1, *Nitrososphaera*
39
40 355 subcluster 4 and *Nitrososphaera* subcluster 8. These subclusters were described previously
41
42 356 by Pester *et al.* (2012). Strong variations in *Nitrosopumilus* subcluster 1 and *Nitrososphaera*
43
44 357 subcluster 8 relative abundances were found between soils from the three Halligen (Figure
45
46 358 5). Relative abundances of *Nitrosopumilus amoA* genes were higher in Nordstrandischmoor
47
48 359 than in Hooge soil samples ($P<0.05$). *Nitrososphaera* subcluster 8 relative abundances were
49
50 360 lower in Nordstrandischmoor than in Hooge or Langeness samples ($P<0.05$).
51
52 361 *Nitrososphaera* subcluster 4 relative abundance did not vary significantly between soils
53
54 362 (Figure 5). Besides the dominant subclusters, the *Nitrosopumilus* subcluster 2, the
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56 363 *Nitrososphaera* subclusters 5 and 6 as well as the *Nitrososphaera* sister cluster were

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3 364 detected in low relative abundance (Figure S1 and S2). The overall AOA community
4
5 365 composition differed most strongly between Hooge and Nordstrandischmoor soil samples
6
7 366 (Figure 6).

8
9
10 367 AOB were characterised by pyrosequencing of 16S rRNA gene amplicons using primers
11
12 368 allowing amplification of both betaproteobacterial and gammaproteobacterial AOB genes. A
13
14 369 total of 302,395 raw sequences was obtained, 260,868 sequences of which were of suitable
15
16 370 quality, providing 6,400 high-quality sequences per sample after normalisation (Table S3).
17
18 371 Rarefaction curves were not saturated (Figure S3), but taxonomic analysis of OTUs revealed
19
20 372 that the dominant bacterial phyla typically occurring in soil (*Acidobacteria*, *Actinobacteria*,
21
22 373 *Proteobacteria*, *Verrucomicrobia*, *Bacteroidetes*, *Firmicutes*, *Chloroflexi*, *Planctomycetes*,
23
24 374 and *Gemmatimonadetes*) (Janssen, 2006) were detected. The 16S rRNA gene sequences
25
26 375 comprised representatives of the betaproteobacterial genera *Nitrosospira* and *Nitrosomonas*
27
28 376 (Figure 5). At some Hooge and Langeness sites no sequences belonging to these genera
29
30 377 were found as *Nitrosospira* and *Nitrosomonas* were presumably below the detection limit.
31
32 378 OTUs belonging to a cluster, which only comprises sequences of so far uncultured
33
34 379 *Gammaproteobacteria* were detected (Figure 7). This cluster is phylogenetically closely
35
36 380 related to known gammaproteobacterial AOB (Figure 7). For instance, "OTU_BacUnc_7",
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38 381 which was detected in soil derived from all Halligen, showed the highest similarity to
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40 382 *Nitrosococcus halophilus* strain Nc 4 (93% identity) (Koops et al. 1990) with respect to
41
42 383 cultured bacteria. It is possible that some representatives of the mentioned cluster represent
43
44 384 gammaproteobacterial AOB. Relative abundances of OTUs affiliated to this cluster were
45
46 385 significantly greater in Nordstrandischmoor soil than Hooge or Langeness soils ($P < 0.05$). We
47
48 386 performed PCR using primers *amoA*-3F and *amoB*-4R as described by Purkhold *et al.*
49
50 387 (2000) to detect gammaproteobacterial AOB but no products were obtained. Attempts were
51
52 388 made to culture gammaproteobacterial AOB but after several subculturing steps only
53
54 389 *Nitrosomonadales* 16S rRNA gene sequences were detected in enrichment cultures (Figure
55
56 390 S4). The relative abundances of *Nitrosospira* and *Nitrosomonas* were significantly greater in
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3 391 Nordstrandischmoor soil ($P < 0.05$). *Nitrosomonas* sequences derived from Hallig soils were
4
5 392 closely related to representatives of the *Nitrosomonadaceae* cluster 6b, cluster 9, and the
6
7 393 *Nitrosomonas cryotolerans* lineage. In contrast, *Nitrospira* clusters were not consistently
8
9 394 recovered with treeing methods using 16S rRNA gene sequences to infer relationships
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11 395 (Prosser *et al.* 2014), but *Nitrospira* sequences detected were closely related to those
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13 396 found in salt marsh sediment. Similar to AOA, the composition of bacterial communities
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15 397 differed most strongly between Hooge and Nordstrandischmoor soil samples (Figure 6).
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17

18 398 **Discussion**

19 20 21 399 **AOA outnumber betaproteobacterial AOB in temporarily flooded soil**

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24 400 Estimated *amoA* gene abundances are in the range of those determined for soil or sediment
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26 401 derived from other coastal environments (Moin *et al.* 2009; Marton *et al.* 2015), except that
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28 402 maximum AOA *amoA* gene abundances were approximately one order of magnitude higher
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30 403 in some New England salt marsh sites (Moin *et al.* 2009). Similar to our study, Moin *et al.*
31
32 404 (2009) and Marton *et al.* (2015) also reported greater abundance of AOA than
33
34 405 betaproteobacterial AOB *amoA* genes in New England and Louisiana salt marsh samples.
35
36 406 High abundances of AOA in coastal soil and sediment suggest that these habitats represent
37
38 407 an important niche for archaeal nitrifiers. Nevertheless, AOA do not always outnumber AOB
39
40 408 in marine sediment, salt marshes or river mouths (Bernhard and Bollmann 2010; Li *et al.*
41
42 409 2015; Dini-Andreote *et al.* 2016). Besides AOA and betaproteobacterial AOB, a cluster
43
44 410 phylogenetically closely related to gammaproteobacterial AOB was detected in temporarily
45
46 411 flooded soil based on high-throughput amplicon sequencing of bacterial 16S rRNA genes.
47
48 412 The number of known complete *amoA* sequences from gammaproteobacterial AOB is very
49
50 413 limited (Junier *et al.* 2008), restricting the design of primers targeting total communities of
51
52 414 these microbes in environmental samples. Therefore, we did not perform *amoA*-specific
53
54 415 quantitative PCR to analyse the abundance of gammaproteobacterial AOB, potentially
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56 416 existing in Hallig soil.
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3 417 **AO inhabiting frequently flooded soil tolerate a wide salinity range**
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6 418 The salinity of the Wadden Sea, surrounding the German Halligen, is similar to that of
7
8 419 artificial seawater with a salt concentration of 30‰. Due to frequent flooding,
9
10 420 microorganisms inhabiting Nordstrandischmoor topsoil are often exposed to high salinities.
11
12 421 Accordingly, in contrast to Hooge and Langeness soil slurries, no decrease in AOA and AOB
13
14 422 *amoA* abundance was determined for Nordstrandischmoor soil slurries at the highest salinity
15
16 423 tested (30‰). Salt concentration in Nordstrandischmoor soil changes over time and can be
17
18 424 affected by environmental factors other than seawater inundation (e.g., rain events). Thus,
19
20 425 microorganisms colonising this dynamic habitat must tolerate a wide range of salinities. This
21
22 426 might explain why AO *amoA* abundances in Nordstrandischmoor soil slurries increased or
23
24 427 did not decrease at any of the tested salinities. Due to differences in Hallig-surrounding
25
26 428 protection constructions, Hooge and Langeness are less frequently flooded than
27
28 429 Nordstrandischmoor (Schindler *et al.* 2014a; Schindler *et al.* 2014b). This could be a major
29
30 430 reason why Nordstrandischmoor soil AO show better adaptation toward Wadden Sea water
31
32 431 salinity than those inhabiting Hooge or Langeness soil.
33

34
35 432 Similar to *amoA* abundances, measured ammonia and nitrite/nitrate concentrations indicated
36
37 433 a preference by Hooge and Langeness soil AO for an artificial seawater salinity of 20‰.
38
39 434 Ammonium concentration did not increase in Nordstrandischmoor soil slurries at any of the
40
41 435 salinities investigated. Considering previous studies focusing on coastal soil or sediment, an
42
43 436 inconsistency in relationships between AO abundance and ammonium or nitrite/nitrate
44
45 437 concentration and nitrification potential rate is evident (Dollhopf *et al.* 2005, Marton *et al.*
46
47 438 2015, Moin *et al.* 2009). For instance, no significant correlations were identified at different
48
49 439 New England salt marsh sites between ammonium or nitrate concentrations and potential
50
51 440 nitrification rates, AOA *amoA* abundance, or betaproteobacterial AOB *amoA* abundance
52
53 441 (Moin *et al.* 2009). In contrast, Marton *et al.* (2015) found a significant correlation between
54
55 442 nitrification and betaproteobacterial AOB and AOA abundances when analysing Louisiana
56
57 443 salt marsh samples. Most previous studies of coastal soil or sediment do not assess or
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1
2
3 444 report the presence of gammaproteobacterial AOB. Here, we detected OTUs belonging to a
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5 445 cluster, comprising so far only uncultured bacteria, which is phylogenetically closely related
6
7 446 to known gammaproteobacterial AOB. We could identify this cluster also in the phylogenetic
8
9 447 tree included in the SILVA (Quast et al. 2013) dataset
10
11 448 "SSURef_NR99_128_SILVA_07_09_16_opt.arb". It is possible that representatives of this
12
13 449 cluster contribute to nitrification in Hallig soil and other coastal soils or sediments, potentially
14
15 450 partly explaining the absence of consistent correlations between activity measurements and
16
17 451 AO abundance. Nevertheless, as we could not identify gammaproteobacterial AOB in
18
19 452 enrichment cultures and detected no gammaproteobacterial *amoA* genes, it is also possible
20
21 453 that the cluster comprises no AO.

22 23 24 454 **Selective association of study site and AO phylotype relative abundance**

25
26
27 455 Previous studies reported the absence or low abundance of *Nitrosopumilus amoA* genes at
28
29 456 a number of sites, including woodland, arable, spruce forest, and tundra soil (Nicol *et al.*
30
31 457 2008; Wessén *et al.* 2011; Pester *et al.* 2012), whereas all samples derived from the
32
33 458 German Halligen harboured *Nitrosopumilus amoA*. Different cultured *Nitrosopumilus*
34
35 459 representatives tolerate seawater salinities and preferably grow at near neutral or slightly
36
37 460 alkaline pH (Qin *et al.* 2014). In addition, genes encoding a mechanosensitive channel
38
39 461 protein and ectoine biosynthesis clusters were identified in microbes affiliated to
40
41 462 *Nitrosopumilus* (Walker *et al.* 2010; Blainey *et al.* 2011; Mosier *et al.* 2012).
42
43 463 Mechanosensitive channels protect cells from hypo-osmotic shock when conditions change
44
45 464 from e.g., high osmolarity to low osmolarity, and ectoine, a compatible solute, can be
46
47 465 produced when salinity is high to prevent influx of sea salt into the cell. Frequent flooding of
48
49 466 Nordstrandischmoor soil leads to temporarily high salinity and strong temporal fluctuations,
50
51 467 which potentially explain increased relative abundance of *Nitrosopumilus* in
52
53 468 Nordstrandischmoor samples compared to soil derived from Hooge and Langeness.
54
55 469 *Nitrososphaera* subcluster 8 has been detected in terrestrial environments such as upland
56
57 470 agricultural, pristine forest or meadow soil, and also in estuarine sediment (see consensus

1
2
3 471 tree provided by Pester *et al.* (2012)), but information about ecophysiology of this subcluster
4
5 472 is very limited and no cultured representatives are currently available. High relative
6
7 473 abundances of *Nitrososphaera* subcluster 8 in Hooge and Langeness samples might be due
8
9 474 to soil salinity and pH preferences. Sequences affiliated to *Nitrososphaera* subcluster 4 were
10
11 475 found in various terrestrial and marine habitats such as sandy ecosystem soil, hot springs,
12
13 476 and lake, deep sea or marsh sediment (see consensus tree provided by Pester *et al.*
14
15 477 (2012)). Thus, *Nitrososphaera* subcluster 4 thrives in ecosystems covering a wide range of
16
17 478 environmental conditions, which might explain similar relative abundances of this subcluster
18
19 479 in Hooge, Langeness and Nordstrandischmoor soil.

20
21
22 480 *Nitrosomonadaceae* cluster 6b, cluster 9, and the *Nitrosomonas cryotolerans* lineage
23
24 481 comprise microbes derived from marine habitats and cultured representatives with obligate
25
26 482 salt requirements (optimal NaCl concentration: 300-400 mM) such as *Nitrosomonas marina*
27
28 483 (Koops *et al.* 1991; Prosser *et al.* 2014). It can be assumed that increased abundances of
29
30 484 the identified *Nitrosomonas* and *Nitrospira* representatives in Nordstrandischmoor
31
32 485 samples are associated with high soil salinities. Increased relative abundances of bacteria
33
34 486 belonging to a cluster, which is phylogenetically closely related to known
35
36 487 gammaproteobacterial AOB were detected in Nordstrandischmoor soil. This might be due to
37
38 488 a preference for high salt concentration and pH values. Nevertheless, currently information
39
40 489 on metabolic or physiological features of representatives affiliated to the mentioned cluster is
41
42 490 lacking and no cultures are available. It is possible that isolation of the so far uncultured
43
44 491 *Gammaproteobacteria* might require very specific culturing conditions.

47 **Conclusions**

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49
50 493 AOA outnumbered their betaproteobacterial counterparts in rarely and rather frequently
51
52 494 flooded soil but in contrast to our first hypothesis obvious differences in *amoA* gene
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54 495 abundance between the three islands were not detected. Nevertheless, apart from these
55
56 496 microbial groups gammaproteobacterial AOB might contribute to nitrification in soil
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3 497 temporarily inundated with seawater. AO indigenous to soil naturally exposed to seawater at
4
5 498 a high frequency tolerated a wide range of salinities (e.g., Wadden Sea water salinity),
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7 499 supporting our second hypothesis. Furthermore, the community composition of both AOA
8
9 500 and AOB differed significantly between the islands. In addition to flooding, other factors such
10
11 501 as specific site differences might have evoked the changes in community composition. In
12
13 502 line with our third and fourth hypotheses, increased abundances of archaeal and bacterial
14
15 503 ammonia oxidisers typically occurring in marine environments were detected in soil derived
16
17 504 from the most frequently flooded island. Seawater flooding significantly affects soil salinity
18
19 505 and pH which probably strongly contributed to selection of detected AO in the analysed
20
21 506 samples. Besides salinity and pH other parameters such as oxygen availability, which were
22
23 507 not determined in this study, are influenced by seawater flooding. Based on our data, we
24
25 508 cannot analyse the proportional importance of these parameters with respect to the
26
27 509 distribution of AO in temporarily flooded soil. Nevertheless, our results indicate that potential
28
29 510 future sea level rise and an associated increase in flooding of coastal soils will enrich
30
31 511 nitrifiers affiliated to e.g., *Nitrosopumilus*, comprising many representatives tolerating
32
33 512 seawater salinity and slightly alkaline pH. To test if the changes in AO communities reported
34
35 513 here also occur in other soils temporarily inundated with seawater, further sites representing
36
37 514 gradients of flooding frequency in different geographic regions should be analysed in future
38
39 515 surveys. A cluster comprising uncultured bacteria closely related to gammaproteobacterial
40
41 516 AOB detected in this study has to be studied in more detail. For instance, single cell
42
43 517 sequencing could be used to gain information on the metabolic potential of this cluster.
44
45 518 Currently, no cultured representative from different microbial groups such as *Nitrososphaera*
46
47 519 subcluster 4 inhabiting temporarily flooded soil are available. AO cultures from soil
48
49 520 temporarily flooded by seawater allow investigation of mechanisms enabling tolerance to
50
51 521 e.g., changes in salinity. Genomic analyses of such cultures are required for identification of
52
53 522 genes potentially involved in responses to osmotic stress. In addition, physiological studies
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55 523 are necessary to analyse e.g., tolerance of single ammonia oxidizer species toward changes
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57 524 in salinity.
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56 **Supplementary Material**7
8
9 **Table S1.** Coordinates of the sampling sites.10
11 **Table S2.** Number of amplicon pyrosequencing-derived AOA *amoA* gene sequences and
12
13 OTU-based analyses at 85% genetic similarity.
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1516
17 **Table S3.** Number of amplicon pyrosequencing-derived 16S rRNA gene sequences and
18
19 bacterial diversity as assessed by the Shannon index at 97% genetic similarity.20
21 **Figure S1.** Consensus tree of AOA *amoA* including representative sequences of OTUs
22
23 affiliated to the *Nitrosopumilus* cluster. The following OTUs determined in this study
24
25 represent the *Nitrosopumilus* cluster: OTU1 (Ho, 14.0 ± 3.8%; La: 54.0 ± 9.7%; No: 69.3 ±
26
27 6.8%) and OTU8 (Ho, 0.00000%; La: 0.00004 ± 0.00004%; No: 0.00025 ± 0.00010%).
28
29 Sequences were added to the consensus tree for AOA *amoA* phylogeny constructed by
30
31 Pester *et al.* (2012) as described in the Materials and Methods section. This tree was then
32
33 pruned. Accession numbers and environmental sources are given in square brackets.
34
35 Abbreviations: Ho, Hooge; La, Langeness; No, Nordstrandischmoor. For more details with
36
37 respect to the different subclusters see Pester *et al.* (2012).
38
3940
41 **Figure S2.** Consensus tree of AOA *amoA* including representative sequences of OTUs
42
43 affiliated to the *Nitrososphaera* and *Nitrososphaera* sister cluster. The following OTUs
44
45 determined in this study represent the *Nitrososphaera* cluster: OTU2 (Ho, 52.4 ± 14.1%; La:
46
47 25.3 ± 10.6%; No: 0.03 ± 0.02%), OTU3 (Ho: 32.8 ± 11.2%; La: 20.6 ± 6.5%; No: 29.8 ±
48
49 6.7%), OTU4 (Ho, 0.00116 ± 0.00063%; La: 0.00000%; No: 0.00000%), OTU5 (Ho, 0.00699
50
51 ± 0.00361%; La: 0.00052 ± 0.00042%; No: 0.00801 ± 0.00244%) and OTU6 (Ho, 0.00022 ±
52
53 0.00011%; La: 0.00000%; No: 0.00006 ± 0.00006%). In addition, the *Nitrososphaera* sister
54
55 cluster is represented by OTU7 (Ho, 0.00000%; La: 0.00009 ± 0.00006%; No: 0.00000%).
56
57 Sequences were added to the consensus tree for AOA *amoA* phylogeny constructed by
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2
3 550 Pester *et al.* (2012) as described in the Materials and Methods section. This tree was then
4
5 551 pruned. Accession numbers and environmental sources are given in square brackets.
6
7 552 Abbreviations: Ho, Hooge; La, Langeness; No, Nordstrandischmoor. For more details with
8
9 553 respect to the different subclusters see Pester *et al.* (2012).
10

11 554 **Figure S3.** Rarefaction analyses of bacterial OTUs at 97% genetic similarity in Hooge,
12
13 555 Langeness and Nordstrandischmoor soil.
14

15
16
17 556 **Figure S4.** Relative abundances of dominant bacterial genera in enrichment cultures. Soil
18
19 557 derived from sampling sites Ho1, No1, No2 and No3 was used as inoculum. The values 30%
20
21 558 and 50% indicate the dilution of artificial seawater used for enrichment (see Materials and
22
23 559 Methods section). Enrichment medium exhibited a concentration of 500 μM or 12.5 mM
24
25 560 NH_4Cl .
26

27 28 561 **Acknowledgements**

29
30
31 562 We thank Heather Richmond and Mechthild Bömeke for providing excellent technical
32
33 563 assistance. In addition, we thank Jessica Heublein for support with respect to basic soil
34
35 564 analyses and Laura Lehtovirta-Morley for useful discussion on cultivation of ammonia
36
37 565 oxidisers. We also thank Ruth Hartwig-Kruse, Michael Kliesch and the team of the
38
39 566 "Schutzstation Wattenmeer Langeness" for support during sampling. This work was
40
41 567 financially supported by the Deutsche Forschungsgemeinschaft (DFG) (NA 848/1-1). The
42
43 568 authors declare that no conflict of interest exists.
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3 **Tables**
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7 **Table 1.** Characteristics of soil samples from Hooge, Langeness and Nordstrandischmoor.
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Soil property	Study sites			ANOVA
	Hooge	Langeness	Nordstrandischmoor	<i>P</i>
Electrical conductivity (mS cm ⁻¹)	3.6 ± 0.4 a	4.0 ± 0.3 a	7.6 ± 0.8 b	<0.001
pH	6.5 ± 0.4 a	6.5 ± 0.0 a	7.6 ± 0.1 b	0.009
Total C (g kg ⁻¹)	50.0 ± 8.3 a	48.0 ± 3.0 a	46.5 ± 2.6 a	0.632
Organic C (g kg ⁻¹)	49.8 ± 8.3 a	48.0 ± 3.0 a	45.1 ± 2.7 a	0.547
Total N (g kg ⁻¹)	4.3 ± 0.7 a	3.9 ± 0.2 a	4.1 ± 0.3 a	0.589
NH ₄ ⁺ -N (mg kg ⁻¹)	12.1 ± 2.5 a	5.2 ± 0.3 b	8.2 ± 1.0 ab	<0.001
NO ₃ ⁻ -N (mg kg ⁻¹)	0.6 ± 0.1 ab	0.4 ± 0.0 a	1.1 ± 0.6 b	0.022

20 Data are presented as mean and standard errors. Thirty soil samples, described in the
21 Materials and Methods section, were considered. Significant ANOVA *P* values are shown in
22 bold (*P* < 0.05). Figures followed by different letters indicate differences among Halligen.
23 NO₂⁻ concentration was negligible and NO_x is therefore equivalent to NO₃⁻-N mg kg⁻¹.
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Figure legends

Figure 1. Maps depicting the German Halligen Hooge (Ho), Langeness (La), and Nordstrandischmoor (No). The locations of the three Halligen and the ten sampling sites are highlighted by black arrows and coloured crosses, respectively. Triplicate samples, 1 m apart, were taken from each site. The large scale map was generated using the Ocean Data View software version 4.7.10 (Schlitzer 2017) (<http://odv.awi.de/>).

Figure 2. Abundance of AOA and betaproteobacterial AOB *amoA* genes in Hooge, Langeness, and Nordstrandischmoor soils samples. Data are presented as mean and standard error of triplicate samples for each site. Abbreviations: Ho, Hooge; La, Langeness; No, Nordstrandischmoor. A total of 30 soil samples was analysed.

Figure 3. Abundance of (a) AOA and (b) betaproteobacterial *amoA* genes in soil slurries incubated for 30 days. Slurries contained soil from Hooge (Ho), Langeness (La), or Nordstrandischmoor (No) and artificial seawater at salinities of 5‰, 10‰, 20‰ or 30‰. Data are presented as mean and standard error of triplicate samples for each site. Asterisks indicate a significant increase (green asterisks) or decrease (red asterisks) in *amoA* abundance after incubation for 15 or 30 days ($P < 0.05$) (the asterisks indicate changes from the initial abundance determined at day 0). A total of 108 soil slurry samples was analysed.

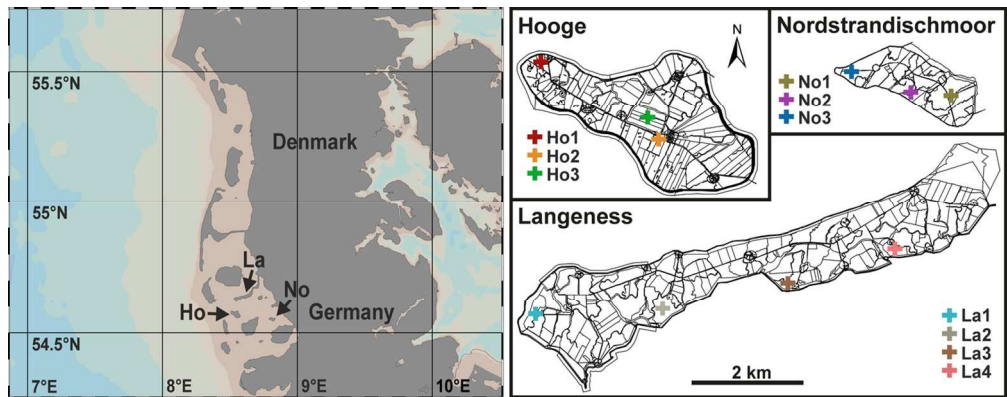
Figure 4. Changes in ammonium and combined nitrite plus nitrate concentrations during incubation of soil slurries for 30 days.

Figure 5. Relative abundances of dominant AOA *amoA* OTUs determined at 85% genetic similarity and *Nitrosospira* and *Nitrosomonas* 16S rRNA gene sequences. Data are presented as mean and standard error. Different letters above the bars represent significant differences in relative abundance of ammonia oxidiser subclusters or genera ($P < 0.05$). Abundances of *Nitrosospira* and *Nitrosomonas* relative to all bacterial sequences are depicted. Abbreviations: Ho, Hooge; La, Langeness; No, Nordstrandischmoor; subcl., subcluster. A total of 30 soil samples were analysed.

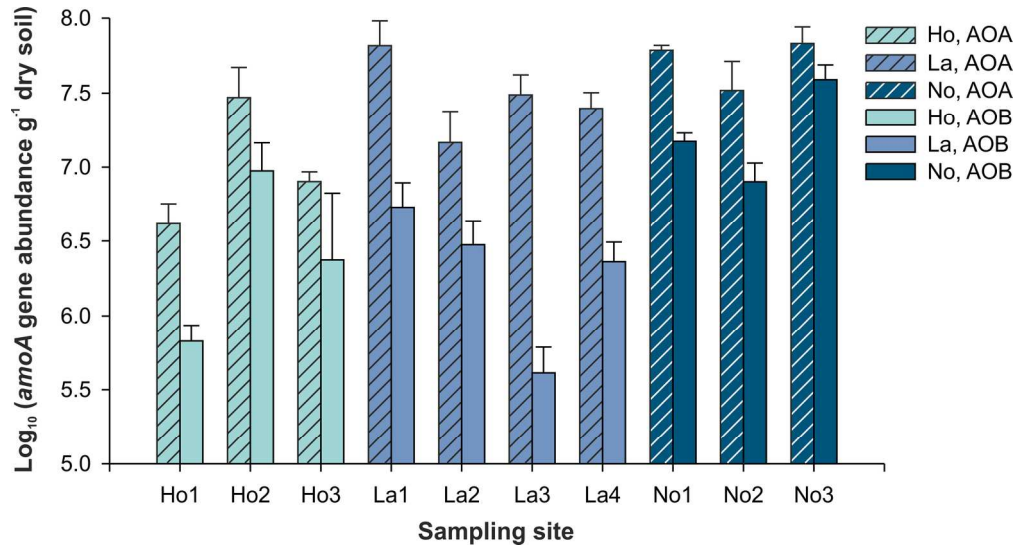
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3 **Figure 6.** Non-metric multidimensional scaling (NMDS) of AOA and bacterial community
4 composition based on weighted Unifrac. Abbreviations: Ho, Hooge; La, Langeness; No,
5 Nordstrandischmoor.
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10 **Figure 7.** Phylogenetic tree based on 16S rRNA gene sequences affiliated to selected OTUs
11 identified in this study, uncultured *Gammaproteobacteria* closely related to these OTUs,
12 *Chromatiaceae*, and *Acidithiobacillus*. Bootstrap values $\geq 50\%$ are indicated at the
13 corresponding nodes. The bar represents 0.02 changes per nucleotide position. Abundances
14 of OTUs relative to all bacterial sequences are depicted for Hooge (Ho), Langeness (La) and
15 Nordstrandischmoor (No) soil. Some of the considered OTUs could only be detected in soil
16 derived from one or two Halligen. Accession numbers and environmental sources are given
17 in square brackets. The 16S rRNA gene sequence of *Nitrosococcus wardiae* D1FHS (Wang
18 *et al.* 2017) has been provided by Drs. Lin Wang and Martin G. Klotz, UNC Charlotte.
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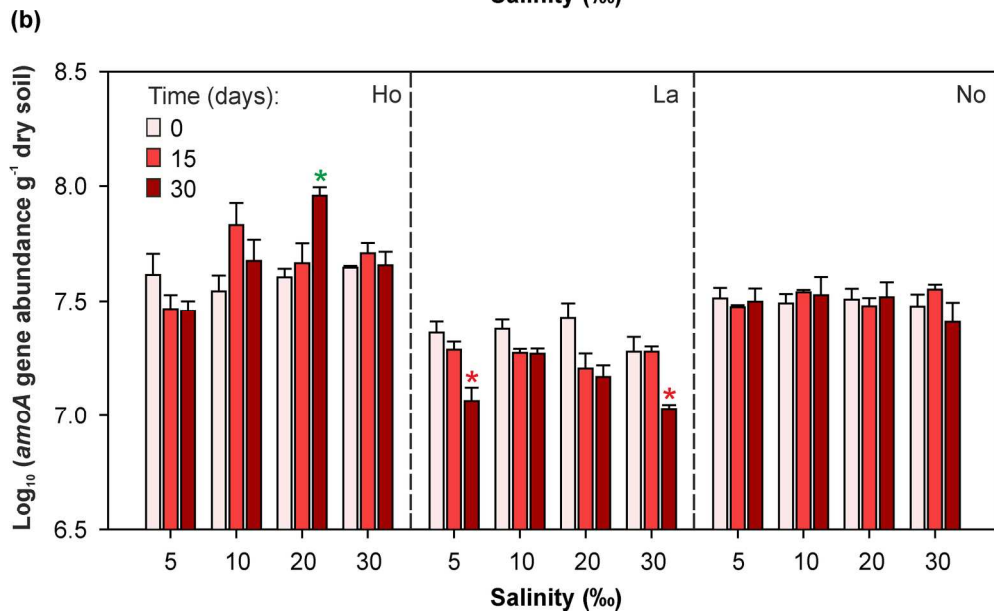
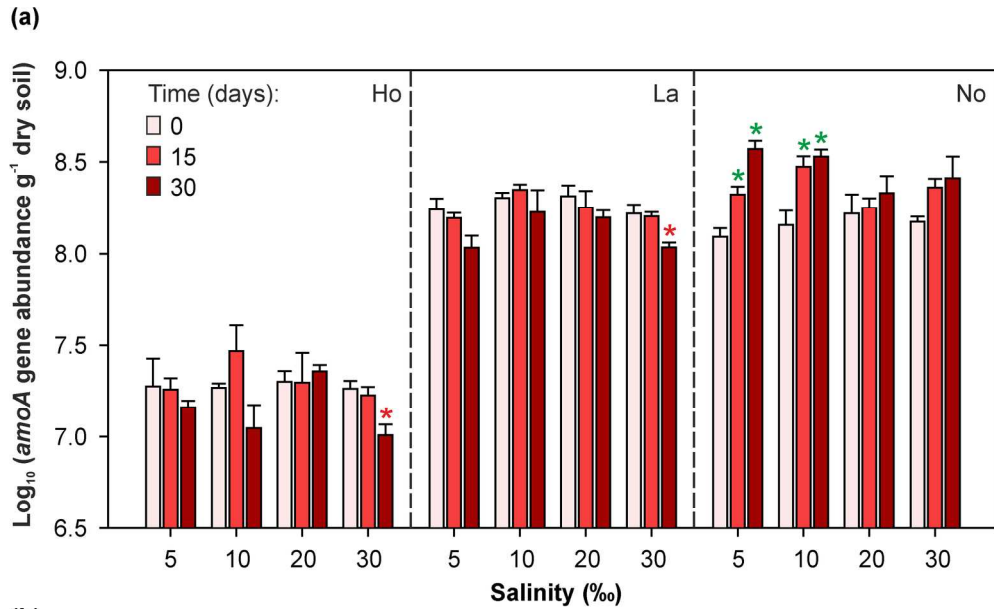


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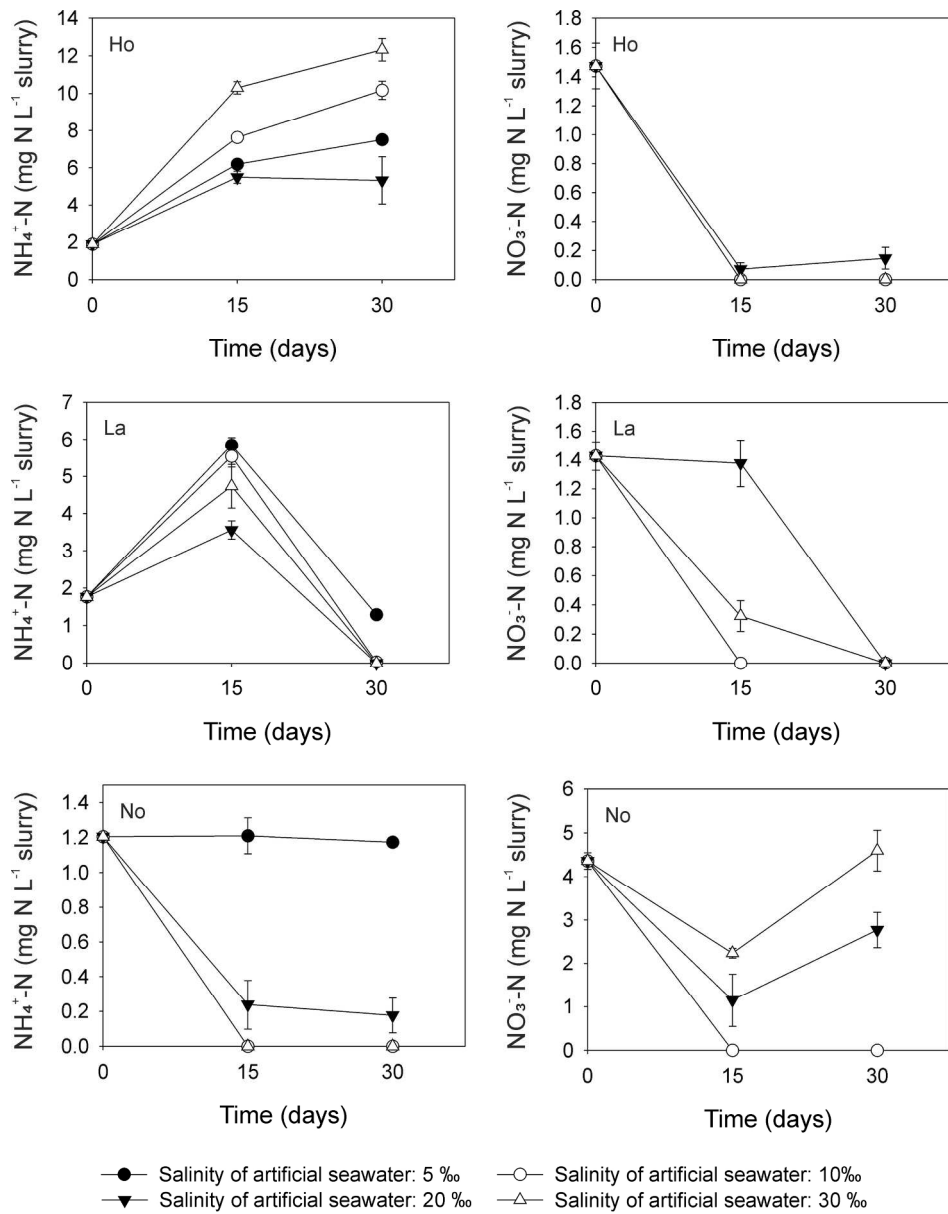
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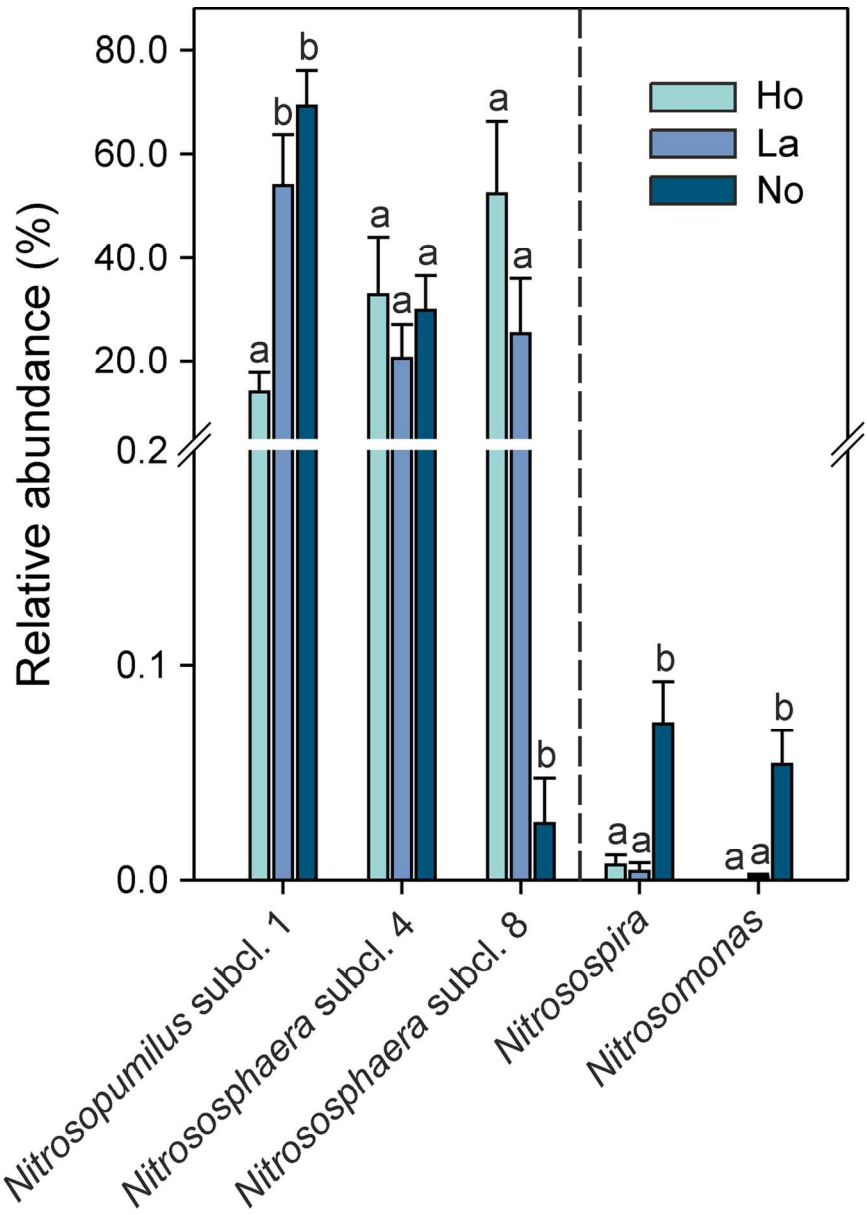


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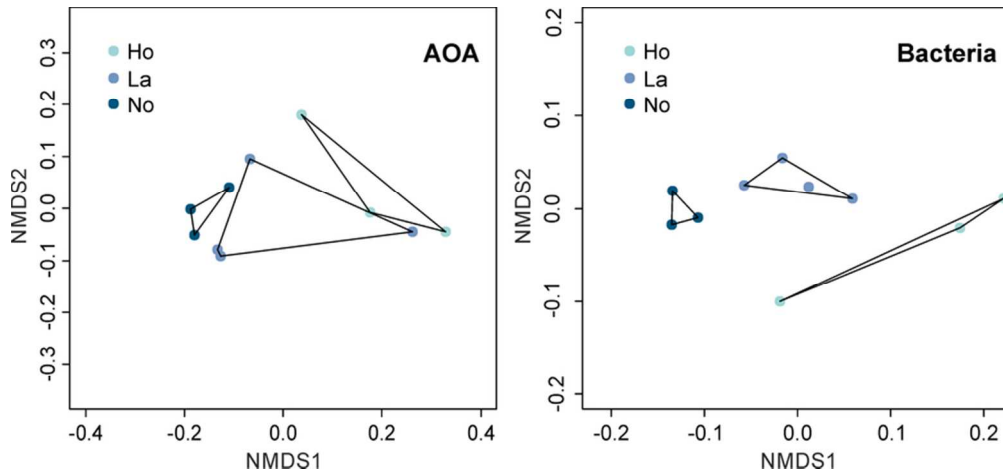


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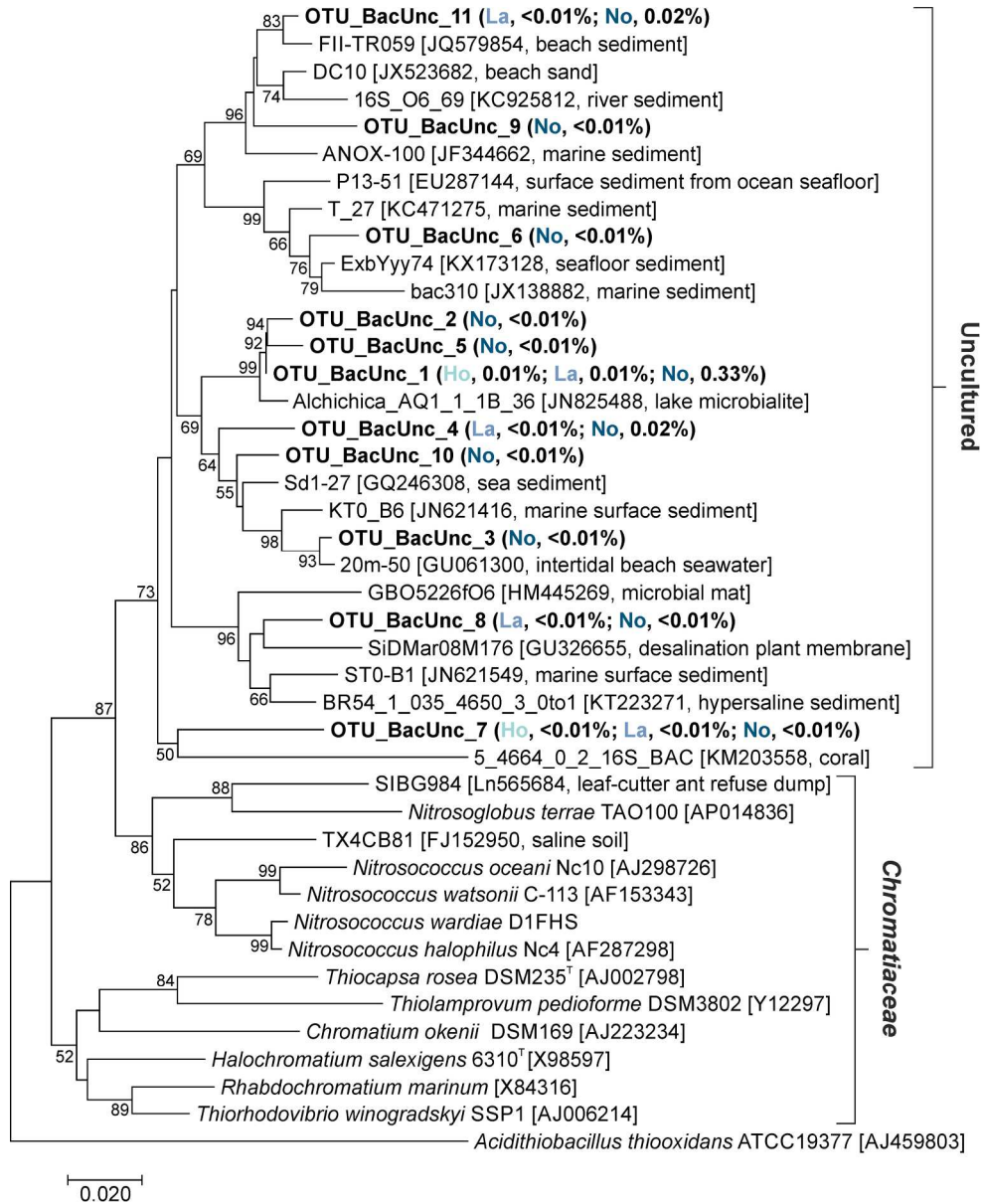


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Table S1. Coordinates of the sampling sites.

Hallig	Sampling site	Latitude (°N)	Longitude (°E)
Hooge	Ho1	54.57817	8.51445
Hooge	Ho2	54.56572	8.54817
Hooge	Ho3	54.56888	8.54468
Langeness	La1	54.62908	8.53368
Langeness	La2	54.63002	8.57032
Langeness	La3	54.63443	8.60470
Langeness	La4	54.63967	8.63360
Nordstrandischmoor	No1	54.54867	8.82168
Nordstrandischmoor	No2	54.54857	8.80983
Nordstrandischmoor	No3	54.55162	8.79447

Table S2. Number of amplicon pyrosequencing-derived AOA *amoA* gene sequences and OTU-based analyses at 85% genetic similarity.

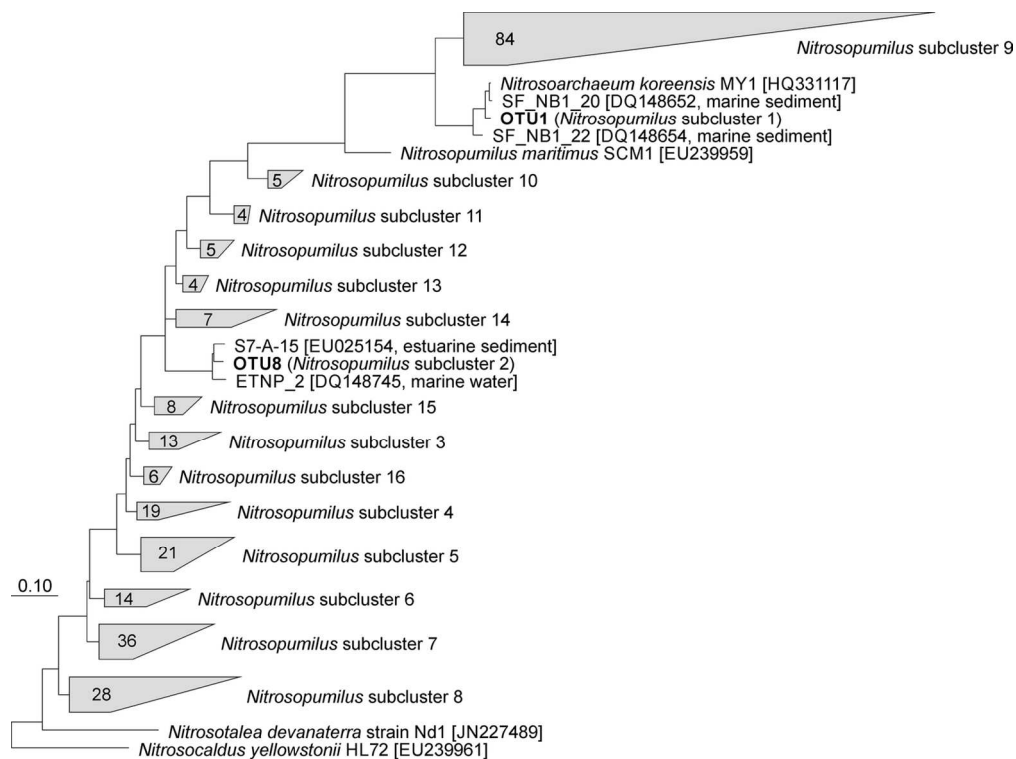
Sample	Number of raw sequences	Number of pre-processed sequences	Number of observed OTUs	Chao1-based OTU number prediction	Shannon index
Ho1a	2,795	1,760	4	4	0.72
Ho1b	2,429	1,512	5	5	0.56
Ho1c	2,598	1,428	5	5	0.64
Ho2a	2,550	1,945	4	4	0.07
Ho2b	2,437	1,788	4	4	0.20
Ho2c	2,125	1,646	3	3	0.03
Ho3a	2,374	1,728	5	5	0.98
Ho3b	2,326	1,777	5	6	0.97
Ho3c	2,261	1,729	5	6	0.98
La1a	2,886	2,023	5	5	0.60
La1b	2,282	1,540	3	3	0.78
La1c	2,458	1,781	3	3	0.80
La2a	2,686	1,953	5	5	0.56
La2b	2,550	1,766	4	4	0.32
La2c	2,493	1,863	3	3	0.93
La3a	2,627	2,055	3	3	0.47
La3b	2,748	2,053	3	3	0.65
La3c	1,697	1,218	3	3	0.33
La4a	2,594	1,704	4	4	0.24
La4b	2,775	1,847	3	3	0.40
La4c	2,706	1,890	3	3	0.78
No1a	2,269	1,588	5	6	0.51
No1b	2,497	1,868	3	3	0.64
No1c	2,775	1,766	4	4	0.56
No2a	2,801	2,055	4	4	0.46
No2b	2,572	1,864	3	3	0.61
No2c	2,789	2,015	5	5	0.58
No3a	2,847	2,011	3	3	0.68
No3b	3,107	2,131	3	3	0.75
No3c	2,373	1,818	4	4	0.31
All samples	76,427	54,122	8	8	1.08

OTU analysis of single samples was performed with the same number of sequences (1,200 randomly selected high-quality sequences per sample).

Table S3. Number of amplicon pyrosequencing-derived 16S rRNA gene sequences and bacterial diversity as assessed by the Shannon index at 97% genetic similarity.

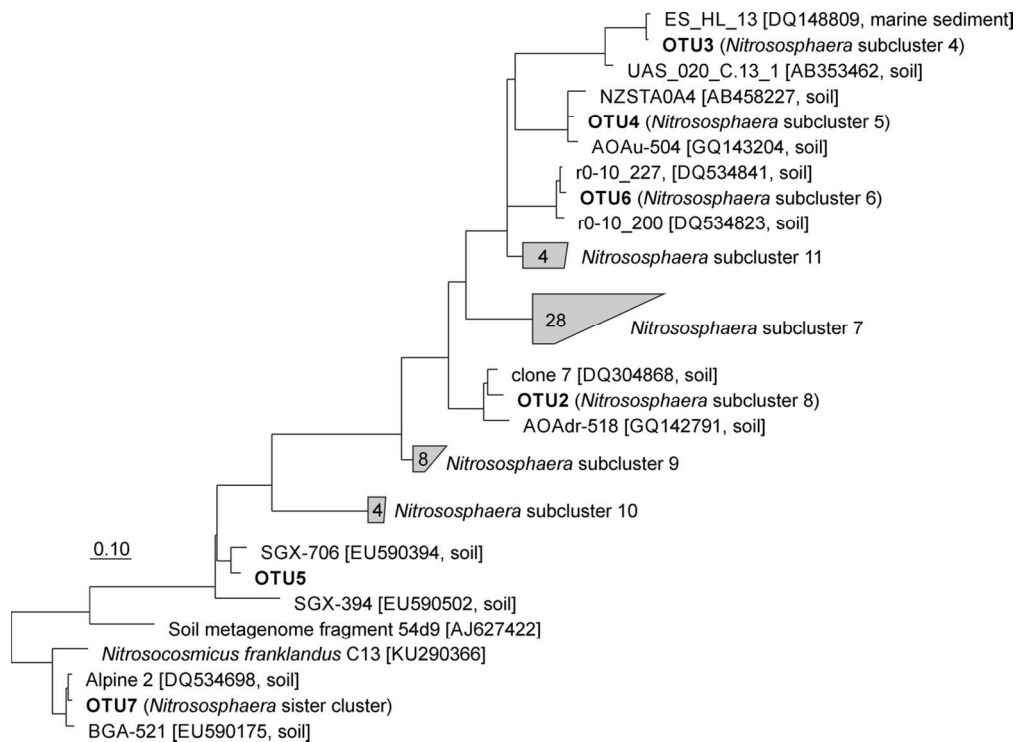
Sample	Number of raw sequences	Number of preprocessed sequences	Shannon index
Ho1a	9,628	8,418	6.39
Ho1b	12,062	10,285	6.54
Ho1c	7,310	6,408	6.56
Ho2a	10,104	8,797	5.65
Ho2b	10,402	9,155	5.62
Ho2c	10,015	8,329	5.50
Ho3a	12,550	11,057	5.57
Ho3b	10,805	9,386	5.40
Ho3c	9,788	8,732	5.45
La1a	12,400	10,483	6.15
La1b	8,347	6,977	6.31
La1c	9,104	7,526	6.42
La2a	10,793	9,383	6.09
La2b	9,386	8,285	6.31
La2c	9,752	8,577	6.10
La3a	8,926	7,623	5.92
La3b	7,457	6,526	5.90
La3c	8,771	7,384	6.07
La4a	9,748	8,300	6.16
La4b	10,668	9,246	6.00
La4c	8,248	7,303	6.03
No1a	11,541	10,092	6.24
No1b	11,021	9,173	6.12
No1c	10,797	9,078	6.15
No2a	10,836	9,474	6.26
No2b	10,386	8,704	6.40
No2c	9,963	8,698	6.30
No3a	10,472	9,082	6.30
No3b	9,984	8,701	6.37
No3c	11,131	9,686	6.32

Calculation of the Shannon index was performed with the same number of sequences (6,400 randomly selected high-quality sequences per sample).

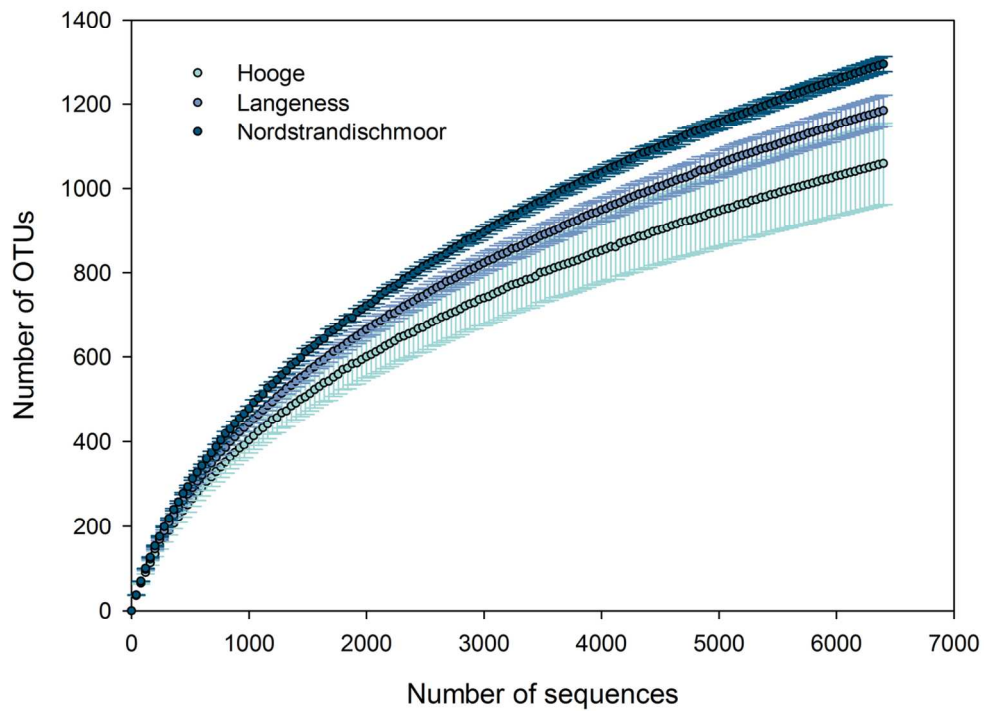


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review

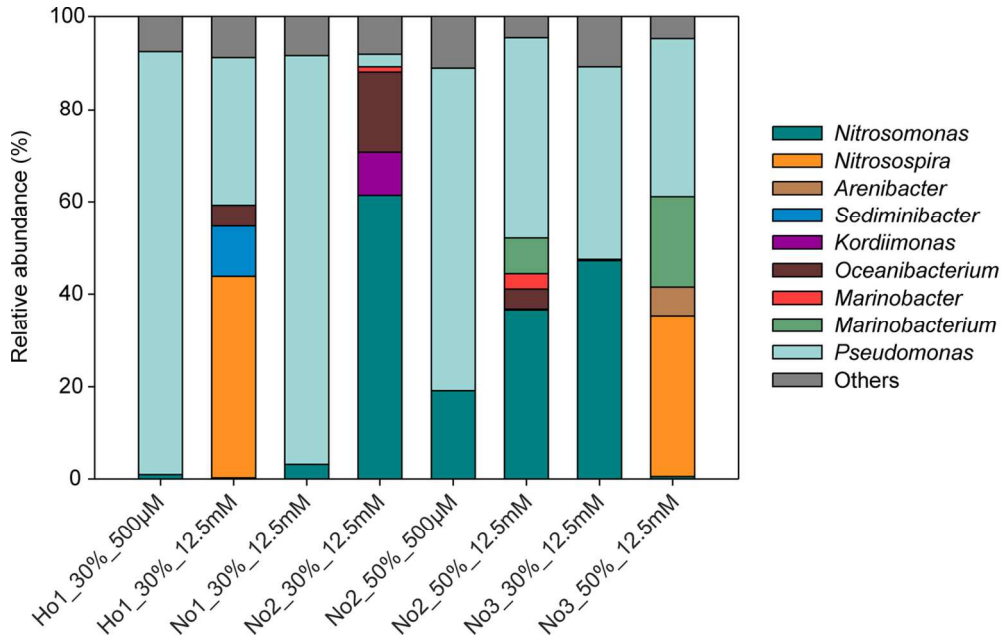


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Review



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Review