Information footprint of different ecohydrological data sources using multi-objective calibration of a physically-based model as hypothesis testing





a-system patterns

Potential insights

Disentangle feedbacks/non-linearities Behaviour outside recorded conditions

Useful information content? Observation Behavour model datasets Figure 1. Opportunities and issues

Nydrologica

Processes

Physically-based

Fully-distributed

Σ

Numerous

parameters

Critical zone science seeks an integrated understanding of hydrological processes considering ecological, geological, geomorphological and pedological couplings [1]. Physically-based, fully-distributed modelling is a promising approach, but its inherent complexity (parameterization) requires a consistent identification of feasible model configurations. Can we use ecohydrological datasets across processes to bring differentiated information content to such models? We tested this approach using the <u>EcH₂O model</u> in a long-term monitored <u>high-latitude catchment</u> [2], where ecohydrological couplings are poorly understood and potentially climate-sensitive.

raised by process-based, fullydistributed modelling approaches in ecohydrology.

Main findings

- Ecohydrological model captures multi-process response in wet, energy-limited steep catchment.
- Using all observation types (streamflow, soil moisture, pine transpiration and net radiation) for calibration yields best performance and lowest predictive uncertainty (Fig. 2).
- Stream discharge brings poorly-focused leverage to ecohydrological simulations.
- Riparian soil moisture and transpiration observations are most informative.

Calibration RMSE *											PI I*												
constraint						0 0.5 1													1	2	2		
Discharge	0.68	0.11	0.5	0.28	0.36	0.31	0.85	1.07	0.96	0.71	1.01	1.08	0.93	0.66	0.48	0.41	0.5	0.71	0.66	0.57	2.75	2.67	2
$ heta_{peat}$	0.65	0.03	0.28	0.23	0.32	0.21	0.96	1.19	1.08	0.71	0.99	1.08	0.93	0.65	1.47	0.04	0.71	0.6	0.59	0.49	3.3	3.21	3
θ_{gley}	0.63	0.19	0.05	0.31	0.41	0.24	0.93	1.21	1.07	0.71	1	1.09	0.93	0.65	1.5	0.68	0.05	0.73	0.74	0.55	3.03	3.05	3
$\theta_{\rm podzol}$	0.65	0.22	0.27	0.06	0.17	0.18	0.95	1.08	1.02	0.71	1.01	1.11	0.94	0.62	1.18	0.65	0.67	0.09	0.09	0.37	3.21	3.01	3
$\theta_{\rm ForestB}$	0.69	0.2	0.31	0.11	0.09	0.18	0.96	1.19	1.07	0.72	1	1.09	0.94	0.64	1.45	0.65	0.72	0.22	0.14	0.43	3.22	3.08	3
$ heta_{all}$	0.59	0.05	0.07	0.1	0.1	0.08	1.01	1.24	1.13	0.71	1.01	1.09	0.93	0.6	1.23	0.14	0.17	0.18	0.17	0.16	2.78	3.18	2
<i>Tp</i> ForestA	0.64	0.15	0.27	0.28	0.3	0.25	0.33	0.7	0.51	0.71	1.01	1.07	0.93	0.55	1.33	0.55	0.6	0.83	0.75	0.69	0.64	0.96	; (
<i>Tp</i> ForestB	0.62	0.16	0.3	0.28	0.37	0.27	0.55	0.4	0.47	0.71	1	1.04	0.92	0.54	1.42	0.51	0.7	0.74	0.71	0.67	0.77	0.63	; (
Tp _{all}	0.62	0.16	0.38	0.26	0.3	0.27	0.37	0.5	0.43	0.71	1.01	1.08	0.93	0.54	1.25	0.53	0.65	0.71	0.63	0.63	0.71	0.79	0
Rn _{valley}	0.62	0.16	0.31	0.27	0.38	0.28	1.02	1.38	1.2	0.68	1.08	1.08	0.95	0.7	1.17	0.48	0.68	0.73	0.71	0.65	3.18	3.09	3
Rn _{bog}	0.6	0.16	0.37	0.33	0.44	0.33	1.04	1.39	1.22	0.72	0.98	1.04	0.91	0.71	1.21	0.39	0.69	0.75	0.75	0.65	3.51	3.38	3
Rn _{hilltop}	0.65	0.25	0.37	0.26	0.33	0.3	0.95	1.27	1.11	0.71	1.01	0.88	0.86	0.67	1.06	0.75	0.76	0.69	0.66	0.71	3.03	3	3
Rn _{all}	0.64	0.21	0.43	0.24	0.3	0.29	0.97	1.27	1.12	0.67	1.02	0.92	0.87	0.67	1.39	0.61	0.6	0.65	0.65	0.62	3.17	3.47	3
All	0.53	0.09	0.16	0.12	0.16	0.13	0.64	0.88	0.76	0.69	1	0.95	0.88	0.52	0.83	0.27	0.48	0.32	0.29	0.34	1.8	1.93	; 1
Simulated variable	Oj.	Openaro Charo		O PO	OK OF	O all	, PX	NOX COLOSU	A Porest	PAR	Philley	Punod	Phillipological	A11	Oj.	Operato Charge		BROU V	0×01	Dall Resth	YOX.	XQX Solosy	10, 10

Figure 2. Heat map of best-runs-averaged (a) root mean square error between model and observations, normalized with the observation average (RMSE*_{m.o}), and **(b)** normalized predictive uncertainty (PU*): daily 90%-spread interval across the 30 best runs divided by the inter-run mean, then averaged over the evaluation period. The x-axis gives the variable or group of variables evaluated, the y-axis shows the dataset(s) used as a constraint over the calibration period.

Outlook

- <u>Robust basis for water pathways characterization</u> across ecohydrological compartments: \rightarrow using process-based tracking of stable isotopes and water age (Kuppel et al., *in prep.*) \rightarrow prediction of consequences from land use and climate alterations
- <u>Advocates for diversifying observations in catchment instrumentation (when possible) for</u> advancing mechanistic understanding of critical zone functioning
- Is this model-data approach efficient for other climatic and topographical settings? \rightarrow Planned application to other well-instrumented catchments across the wider north







1 0.4 0.4 0.59 0.46 0.96 <mark>6</mark> 0.3 0.3 0.58 0.39 1.1 4 0.37 0.37 0.6 0.45 1.1 0.36 0.36 0.68 0.47 1.03 0.41 0.43 0.64 0.49 1. 0.35 0.38 0.6 0.44 <mark>0.92</mark> .8 0.34 0.36 0.69 0.46 0.7 7 0.36 0.4 0.57 0.44 0.68 <mark>5</mark> 0.36 0.39 0.62 0.46 0.6 4 0.07 0.06 0.64 0.25 **1.08** <mark>5</mark> 0.21 0.17 0.56 0.31 <mark>1.</mark>1 0.33 0.36 <mark>0.03</mark> 0.24 1.06).11 0.13 0.12 0.12 <mark>1.0</mark>9 15 0.19 0.25 0.2 <mark>0</mark>.6 P3 4/1 P

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Figure 3. Bruntland Burn catchment, showing (a) topography, stream network, and micrometeorology (including net radiation, star). The conceptualisation used for simulations (b-f, 30×30m² resolution) comprises (b) pedology, aggregated from the Hydrology of Soil Types (HOST) classes, and (c-f) the pixel fraction covered by the four considered vegetation types (in addition to scree/bare soil, not shown).



Chris Soulsby ^{1,2}, Sylvain Kuppel ¹, Marco P. Maneta ³ and Doerthe Tetzlaff ^{1,2,4}

¹Northern Rivers Institute, University of Aberdeen, United Kingdom. ²Leibniz Institute of Freshwater Ecology and Inland Fisheries, Germany. ³ Geosciences Department, University of Montana, USA. ⁴ Department of Geography, Humboldt University Berlin, Germany.

- [6] Ali, G., et al., Earth Surf. Process. Landf. **39**, 399–413 (2014).