- 1 **Title:** Late Pleistocene prey mobility in southwestern France and its implications for
- 2 reconstructing Neandertal ranging behaviors
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- 4 Authors: Jamie Hodgkins^{a, 1,*}, Alex Bertacchi^b, Kelly J. Knudson^c, Troy Rasbury^d, Julia I.
- 5 Giblin^e, Gwyneth Gordon^f, Ariel Anbar^{f,g}, Alain Turq^h, Dennis Sandgatheⁱ, Hannah M.
- 6 Keller^b, Kate Britton^j, Shannon P. McPherron^k
- 7

8 Affiliations:

^a Department of Anthropology, University of Colorado Denver, 1200 Larimer Street Denver, CO
 80217-3364, USA

- ^b Department of Anthropology, Yale University, New Haven, CT 06511, USA
- ^c Center for Bioarchaeological Research, School of Human Evolution and Social Change,
 Tempe, Az, 85287-2402, USA
- ^d Department of Geosciences, Stony Brook University, Earth and Space Science Building, Stony
 Brook, NY 11794-2100, USA
- ^e Department of Sociology and Anthropology, Quinnipiac University, Hamden, CT 06518, USA
- ¹⁷ ^f School of Earth and Space Exploration, Arizona State University, Tempe, AZ 85287, USA
- ^g Department of Chemistry and Biochemistry, Arizona State University, Tempe, AZ 85287, USA
- 19 ^h Museé National de Préhistoire, Les Eyzies 24200, FRANCE
- ⁱ Department of Archaeology, Simon Fraser University, Burnaby, BC V5A-1S6, CANADA; and,
- University of Pennsylvania Museum of Archaeology and Anthropology, Philadelphia, PA
 19104, USA
- ^j Department of Archaeology, School of Geosciences, University of Aberdeen, Elphinstone
 Road, Aberdeen, AB24 3UF, Scotland
- ^k Department of Human Origins, Max Planck Institute for Evolutionary Anthropology, Deutscher
 Platz 6, 04103 Leipzig, Germany
- 27
- 28
- 29 *Correspondence to: Jamie.Hodgkins@ucdenver.edu, Phone: 720-217-9253
- 30 ¹permanent address: Department of Anthropology, University of Colorado Denver, 1200 Larimer
- 31 Street, Denver, CO 80217-3364, USA
- 32
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34 Highlights

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- We use strontium isotope data from two Late Pleistocene (Middle Palaeolithic) sites
 in southwestern France to explore the movement behaviors of large ungulates
- Our results suggest that the main prey species (reindeer and bison) did not
 undertake movements across the Massif Central and that their range use did not
 extend beyond the basins of France.

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43 Key Words

- 44 ⁸⁷Sr/⁸⁶Sr analysis, Neandertal prey mobility, reindeer migration, Aquitaine, Dordogne, Pech de
- 45 l'Azé IV, Roc de Marsal, paleoecology

47 ABSTRACT

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49 As hunter-gatherers, Neandertal mobility and corresponding adaptations were influenced 50 by the mobility of their prey; thus, it is critical to track how the movement patterns of each 51 species varied over time at specific sites. Here, prey paleomobility is reconstructed by measuring radiogenic strontium isotope ratios (⁸⁷Sr/⁸⁶Sr) in herbivore teeth recovered from two 52 53 archaeological sites (Pech de l'Azé IV and Roc de Marsal) in the Dordogne Valley (Aquitaine 54 Basin) of southwestern France that span marine isotope stages (MIS) 5-3. These ratios are 55 compared to a published isoscape as well as new environmental samples (soil, *Helix* shells, 56 rodent teeth, water, and wine). Our results show that reindeer (*Rangifer*) likely moved within the 57 sedimentary basins of France and did not venture into mountainous regions. Instead, reindeer, 58 bison (Bison) and other prey species were available year-round in the Aquitaine and Paris basins. 59 Red deer (*Cervus*) and horse (*Equus*) may have ventured into the mountains. These results 60 suggest that Neandertals could find year-round faunal resources in these basins, at least during 61 MIS 4 and 3.

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64 **1. Introduction**

In interpreting Neandertal hunting behaviors, it is often thought that European Pleistocene ungulates, particularly reindeer (*Rangifer tarandus*), undertook long distance round-trip migrations (Bahn, 1977; Burch, 1972; Gordon, 1988; Morin, 2004; Straus, 1997; Sturdy, 1975). For example, zooarchaeological studies in France proposed *Rangifer* spent their winters in southwestern France and northern Spain, but during the summer they would move to areas with tundra (e.g., Scandinavia or Switzerland), or up into the mountains close to glaciers (e.g., the Massif Central or Pyrénées), or that they would make seasonal movements between the Massif 72 Central and the Bay of Biscay (Arbos, 1932; Bahn, 1977; Dibble and Rolland, 1992; Gordon, 1988; Mellars, 1985; Straus, 1997). Archaeological evidence supporting the hypothesis that 73 74 *Rangifer* in southern France were migratory came from seasonality studies based on the eruption, 75 wear, and cementum growth of herbivore teeth and on antler growth patterns (Bahn, 1977; Burke 76 and Pike-Tay, 1997; Delpech, 1983; Enloe and David, 1997; Gordon, 1988). Limited isotope 77 studies undertaken in the region to date also provided some evidence of seasonal migration, but -78 given the lack of diversity in local bioavailable strontium – distances of such migrations have been 79 difficult to infer. Britton et al. (2011) conducted a radiogenic strontium (Sr) isotopic study of three 80 Rangifer and a single Bison individual from the Quina Mousterian unit W-22 of Jonzac (also 81 known as Chez-Pinaud), located in the Charente-Maritime region of southwestern France, and 82 concluded that *Rangifer* from that site were likely migratory (i.e., demonstrated intra-tooth variation in strontium consistent with seasonal movements). However, given that all ⁸⁷Sr/⁸⁶Sr 83 84 ranges fell within variability found within the broader Aquitaine and Paris basins, distances could 85 not be estimated and indeed seasonal movements may have been restricted to those regions with 86 no evidence of traversing mountainous areas. The single Bison specimen from Jonzac fell within 87 local variability immediate to the site and, in contrast to the reindeer, showed no seasonal 88 variability. A radiogenic Sr study of the Middle Paleolithic fauna from the nearby Les Pradelles 89 also did not find evidence of basin-mountain migration (Kelly, 2007). Furthermore, some analyses 90 of antler growth patterns, tooth wear and other palaeoecological studies based on fauna have 91 argued that *Rangifer* in this region were not migrating long distances (see Bouchud, 1966; Mellars, 92 1985), with some authors suggesting that at least some groups of *Rangifer* were present year-round 93 (Bouchud, 1966; Delpech, 1983; Lartet and Christy, 1875; Pike-Tay, 1991; White, 1985). 94 Evidence suggesting that *Rangifer* did not move into the mountains within and surrounding France

and that *Bison* ranged locally is intriguing, and of course mobility patterns of prey species can be
variable by location and time, so it is important to track these movements site by site and through
different climatic phases.

98 In a broad analysis including stone tool reduction, faunal remains and paleoclimatic 99 reconstructions, Dibble and Rolland (1992) suggested that during cold glacial cycles when 100 Rangifer dominated faunal assemblages Neandertals may have had long winter occupations in 101 southwestern France, moving little during the winter months. During interglacial cycles when 102 Bison, Cervus, and Equus dominated archaeological sites Neandertals may have been more mobile, 103 occupying sites for short periods of time. Other researchers have pointed out that the abundance 104 of animal species present during the Middle and Upper Paleolithic of southwestern France may 105 have created a stable and predictable supply of food year-round, possibly leading to less mobile 106 hunter-gatherer groups (Mellars, 1985; Spiess, 1979). Burke and Pike-Tay (1997) have suggested 107 that if seasonality data on Cervus, Rangifer and Equus remains from Late Upper Paleolithic sites 108 in southwestern France are compared to each other, the data point to an annual hunting range of 109 only 40-50 km within the northeastern portion of the Aquitaine basin. If hunter-gatherers living in 110 the Aquitaine basin had access to reliable resources year-round, they may have sustained higher 111 population concentrations (Mellars, 1985; Spiess, 1979), although ethnographic studies of groups 112 whose hunting economy is concentrated on many different terrestrial herbivores have been 113 reported to live in small family groupings (Speck and Eiseley, 1942).

With the goal of exploring the paleoecology of the prey-species Neandertals depended on in southwestern France, thus contextualizing Neandertal behavior, here we use radiogenic strontium isotope analysis of *Rangifer*, *Bison*, *Equus*, *Cervus* and *Capreolus* (roe deer) teeth excavated from two French Neandertal sites: Pech de l'Azé IV (Bordes' 1970-1977 excavation)

and Roc de Marsal (the most recent excavations; hereafter Pech IV and RDM, respectively) in the Dordogne Valley (Aquitaine Basin, SW France) to explore whether or not these species' ranging habits were extending beyond the sedimentary basins of France at that time. We predict that if herbivores were spending time in the mountains, the ⁸⁷Sr/⁸⁶Sr in their teeth should range from relatively low values consistent with foraging in the sedimentary basins of France, to relatively high values consistent with foraging in areas with older geological substrates.

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125 **2. Background**

126 Understanding Neandertal behavior and local ecology requires knowledge of the behaviors 127 of their prey, including whether animals were long distance migrants or stayed within smaller 128 ranges. Relying on supposedly species-typical behaviors to infer prey mobility can be problematic 129 when there is environmentally-mediated variation in how animals move. For example, populations 130 of ungulates that live in more wooded environments migrate less than ungulates that live in tundra 131 and grassland environments because food resources are distributed in patches that are 132 geographically closer together in woodland environments than in open tundra and grassland 133 ecosystems (Johnson et al., 2002a, 2002b, 2001; Senft et al., 1987; Skogland, 1984; Tyler and 134 Øritsland, 1989). Thus, robust paleobiological reconstruction of Neandertal prey mobility should 135 draw on multiple lines of evidence including environmental proxies and direct or indirect evidence 136 for animal movement (e.g., strontium stable isotopes).

Ecologists have emphasized that the word 'migration' is used to describe many types of movement patterns including (but not restricted to): round trip 'seasonal' travel, one-way movements, (often seasonal) altitudinal shifts, and even nomadism, in which an animal does not follow a predictable route but forages opportunistically (Dingle and Drake, 2007). While these

141 movement behaviors manifest themselves in different ways, ultimately migratory behaviors are a 142 response to ecological stimuli and thus different populations within one taxon can exhibit different 143 behaviors (Dingle and Drake, 2007). Long distance, round-trip migratory behavior in large 144 ungulates has been shown to be selectively advantageous for two main reasons: 1) to enhance 145 temporary access to high quality food in areas that cannot be inhabited year-round (resource 146 tracking) and 2) to reduce risk of predation (Abraham et al., 2022; Fryxell and Sinclair, 1988; 147 Tyler and Øritsland, 1989). Exposure to insects, other parasites and diseases also drive movement 148 patterns in large ungulates (Downes et al., 1986; Grayson and Delpech, 2005; Russell et al., 1993). 149 Some populations that live in productively poor ecosystems compensate by migrating over vast 150 distances seasonally, such as *Rangifer* (caribou = North American reindeer, reindeer = European 151 Rangifer) that live on the open tundra and Bison (Bison bison bison) of North America (e.g., some 152 grassland environments). The average round-trip migration distance measured among three 153 different herds of barren ground caribou is 1673 ±491 km, with the longest distance measured at 154 2500 km (Berger, 2004; Bergerud, 1974; Brown et al., 1986). Similarly, *Bison* that occupy habitats that are low in productivity move within areas of 712 km² - 1240 km² (Berger, 2004; Hanson, 155 156 1984; Reynolds et al., 2003).

157 Non-migratory ecotypes of *Rangifer* (often referred to as 'woodland caribou') live 158 throughout Canada (e.g., north-central British Columbia, northeastern Alberta, Ontario, 159 northwestern Quebec, Labrador) and in the northwestern United States (e.g., Montana, Idaho, and 160 Washington) (Brown et al., 1986; Evans, 1960; Johnson et al., 2002b, 2001; Layser, 1974; Stuart-161 Smith et al., 1997). The non-migratory ecotype can also be found today in the Norwegian arctic 162 archipelago of Svalbard and in Finland (Kojola et al., 2004; Tyler and Øritsland, 1989). In North 163 America, herds of non-migratory *Bison* (often referred to as wood bison) historically inhabited

164 much of the same habitats as non-migratory *Rangifer* with species ranges extending from 165 northwestern Canada into eastern Alaska (Reynolds et al., 2003). In Europe, non-migratory Bison 166 inhabited forests in northern, eastern, and western Europe (Pucek et al., 2004); environments 167 composed of mosaic habitats of pine and spruce woodlands with feather mosses and lichens, 168 shrub/sedge and forb-dominated wetlands, river and lake environments, alpine shrublands and 169 alpine grasslands (Johnson et al., 2002b). In general, the migratory behavior of Rangifer and Bison 170 appear to be regulated by the temporal and spatial variability of resources, particularly lichen, 171 sedges, and grasses (Johnson et al., 2002a, 2001; Larter and Gates, 1991; Tyler and Øritsland, 1989). In southwestern France, relatively positive δ^{13} C values of Late Pleistocene *Rangifer* have 172 173 been measured between 33-15 ka (Drucker et al., 2003) suggesting that lichens were contributing 174 to reindeer diets, although this is not the case for all periods and all regions of France. For example, 175 data from the MIS 3 site of Abri du Maras indicate that lichen consumption was low and that 176 reindeer diets were similar to that of other cervids (Britton et al. 2023).

177 To better understand the diversity of prey animals available to Neandertals in the Dordogne 178 valley over time, three additional species found at Neandertal sites in high abundance are analyzed 179 in this study: horse (Equus ferus), red deer (Cervus elaphus), and roe deer (Capreolus capreolus). 180 Living horse populations have been heavily influenced by domestication, altering modern horse 181 behavior and ecology (Bendrey, 2012). van Asperen (2010) notes that horses from the Middle 182 Pleistocene of northwestern Europe seem to have been less mobile than modern Przewalski's 183 horses (Equus ferus przewalskii) and onagers (Equus hemionus). Ecomorphological and isotopic 184 dietary studies of caballoid horses from the Middle to Late Pleistocene demonstrate that they were 185 grazers in predominantly open forests and grasslands, though some horse assemblages suggest that 186 they could also forage in more closed habitats (Boulbes and Van Asperen, 2019). Middle to Late

Pleistocene horse remains have been found at archaeological sites with paleoenvironmental contexts ranging from open glacial steppes to closed forests (van Asperen, 2010). The strontium isotope analysis conducted here will help clarify how this species behaved in the Pleistocene of southwestern France, as well as adding to information about the availability of Neandertal prey at the sites of RDM and Pech IV.

192 A plethora of research has been conducted on the movement patterns of red deer (Cervus 193 elaphus) (Albon and Langvatn, 1992; Georgii, 1980; Georgii and Schröder, 1983; Peek, 2003; 194 Pike-Tay, 1991). Cervus populations are widely distributed with extant populations ranging 195 throughout western North America, northern Africa, and from western Europe to Siberia (Mahmut 196 et al., 2002; Peek, 2003). Many Cervus populations are known to move from lower altitudes in the 197 winter to higher altitudes in the summer (Albon and Langvatn, 1992) and also from coastal to 198 inland areas. These seasonal shifts prolong access to high quality food resources (Albon and 199 Langvatn, 1992). However, as is the case with the other large herbivores studied here, a great deal 200 of variability has been observed in *Cervus* movement patterns (Albon and Langvatn, 1992; 201 Georgii, 1980; Georgii and Schröder, 1983; Peek, 2003; Schmidt, 1993), with some groups 202 (including populations in the Alps) undertaking long distance migrations, while other groups are 203 non-migratory (Georgii, 1980; Georgii and Schröder, 1983; Peek, 2003). Once again, this 204 variability in migratory behavior and home range size appears to be the product of the availability 205 of food resources, difficulty of travel, cover quality, and the presence of insects (Brazda, 1953; 206 Peek, 2003). One study of Pleistocene red deer focused on MIS 6 levels at Lazaret Cave in France. 207 While the red deer showed strong seasonality of movement in their intra-tooth profiles, spatial 208 assignment indicates that both the summer and winter ranges can be found within 20-30 km of the 209 site, and that the migrations were likely to have been altitudinal (summer in the uplands, winter in

the lowlands) (Barakat et al. 2023). This highlights that while strontium can reveal seasonal movement patterns in ungulates effectively when combined with serial-sampling, especially in regions of highly variable lithology (e.g., southwest France), those movements may not be occurring over long distances. Similarly, in areas of homogenous lithology, long-distance movements may not be clear from intra-tooth strontium profiles alone, thus detailed analyses are needed.

216 Roe deer (*Capreolus capreolus*) are small cervids (15-35 kg) that live in woodlands, 217 heathlands, and grasslands (Lovari et al., 2017; Saïd et al., 2005; Tufto et al., 1996). Capreolus 218 are non-migratory, though snow accumulation has been observed to cause some groups to move 219 short distances seasonally. For example, Ramanzin et al. (2007) tracked radio-collared Capreolus 220 in alpine regions of Italy as they moved 12.0 ± 6.2 km from higher-elevation summer ranges to 221 lower elevation winter ranges. Throughout the year median home range size for males is 16.7 ha 222 with a range of 13-31 ha, and 23 ha for females with a range of 13-44 ha (Lovari et al., 2017). 223 Home range size for the species is influenced by plant biomass and reproductive success with more 224 offspring leading to larger home-ranges for females (Saïd et al., 2005). Because Capreolus is (and 225 likely was) a non-migratory species with small home range size, they can be used a control species: 226 the radiogenic isotope analysis of the teeth of Capreolus will indicate what values we should 227 expect from a local prey species, we can then compare the extent to which isotopic values of other 228 species differ.

Knowledge about the environments in which modern populations live and the movement behaviors undertaken by herds living in different types of environments inform us as to how populations of animals would have behaved in the past. Much of the ecological literature has concluded that, during the Pleistocene, areas of southwestern Europe acted as a refugium for trees

233 (e.g., spruce, pine, oak, elm, beech), faunal species and hominins (Barron et al., 2003; de Beaulieu 234 and Reille, 1992; Huntley and Allen, 2003) during cold climatic oscillations. Reconstructed 235 temperature ranges for southern France during MIS 3 (~59-37 ka) appear to have been similar to 236 temperatures observed in modern boreal forests, inhabited by woodland species today (e.g., 237 modern boreal forest = 5 °C in winter and 17 °C in summer) (Barron et al., 2003; Huntley and 238 Allen, 2003; Pederzani et al., 2021). If these temperature reconstructions are accurate, then 239 mammals living in these environments may have behaved more like extant woodland-dwelling 240 ungulates in undertaking movements within the basins of France (Bouchud, 1966; Delpech, 1983; 241 Lartet and Christy, 1875; Pike-Tay, 1991; White, 1985), effectively creating an environment with 242 dispersed but dense patches of resources (Binford, 1980; Kelly, 1992; Lieberman and Shea, 1994). 243 If this is true, Neandertals would have been able to hunt populations of herbivores year-round 244 within the French basins. In this case a pattern of radiating mobility with sites that served as multi-245 seasonal base camps, as proposed by many researchers (Binford, 1980; Kelly, 1992; Lieberman 246 and Shea, 1994; Turg et al., 2017), may be supported. In contrast, temperature reconstructions 247 based on oxygen isotopic ratios in equid enamel phosphates suggest stronger temperature 248 seasonality during MIS 3 compared to today in the Rhône Valley at the site of Abri du Maras (as 249 extreme as -10 °C in winter and 31 °C in summer) and – in that phase of the site at least – these 250 conditions are associated with a migratory ecotype of *Rangifer tarandus* (Britton et al., 2023).

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252 2.2. Radiogenic isotopic migration studies

Radiogenic Sr isotopic studies have been used to track migration patterns in modern and
archaeological animal and human populations (e.g., Britton et al., 2011; 2023; Baraket et al. 2023;
Copeland et al., 2016; Ericson, 1985; Hodgkins et al., 2020, 2018; Hoppe et al., 1999; Knudson et

256 al., 2004; Price et al., 1994; Scharlotta et al., 2011), including Neandertals (e.g., Moncel et al., 257 2019; Richards et al., 2008). To "ground-truth" the method for use in fossil *Rangifer*, Britton et al. 258 (2009) analyzed radiogenic Sr isotope ratios in the teeth of extant Rangifer belonging to the 259 Western Arctic Herd in Alaska for which the migratory behavior was known. Results from their 260 analysis showed that the ratios found in the second and third molars of each deer broadly matched 261 the geological ratios throughout the migration route demonstrating that this is an effective method 262 for tracking *Rangifer* migrations, an approach which has since also proven valid when using laser 263 ablation (Le Corre et al. 2023).

264 As summarized in Price et al. (2002), radiogenic Sr ratios are useful for tracking migrations 265 because the bedrock of a particular geological region has a characteristic isotopic signature. That 266 is because different rock types have different initial content of calcium (Ca) and potassium (K), which are substituted for by Sr and rubidium (Rb), respectively. Through time, ⁸⁷Rb decays to ⁸⁷Sr, 267 increasing ⁸⁷Sr/⁸⁶Sr values. Thus, old K-rich rocks (e.g., granite) have higher ⁸⁷Sr/⁸⁶Sr values than 268 269 young Ca-rich ones (e.g., limestone) (Bentley, 2006). As bedrock is eroded by water, wind and 270 chemical weathering, Sr enters the hydrological system and is taken up by plants, which are 271 ingested by herbivores. Sr is incorporated into herbivore teeth as a substitute for Ca; consequently, the ⁸⁷Sr/⁸⁶Sr signature of an herbivore's teeth is generally indicative of where the animal was as its 272 dentition developed. Thus, if ⁸⁷Sr/⁸⁶Sr within one region is distinct from adjacent areas, measuring 273 ⁸⁷Sr/⁸⁶Sr at multiple loci across the tooth crown that formed over several months can indicate if 274 275 the animal was moving between those areas (Britton et al., 2009; Lazzerini et al., 2021; Miller et 276 al., 2022). While high-density sampling down the entire length of the crown (either using 277 traditional intra-tooth sampling, or laser ablation) is desirable and likely necessary to observe 278 repeated (i.e., back and forth) movement behaviors, spot sampling at a smaller number of intervals

down the crown should be sufficient to identify any potential major shift in ⁸⁷Sr/⁸⁶Sr during tooth formation. This study will employ such an approach – targeting 2-3 sampling points per tooth for low crowned species such as *Rangifer* and *Cervus* and *Capreolus* (closest to the occlusal surface, closest to the enamel-root junction, at the midpoint of the crown), while targeting up to 8 sampling points per tooth on high crowned teeth from species such as *Equus*, to specifically assess whether the herbivores sampled moved between the sedimentary basins in France (i.e. areas with low ⁸⁷Sr/⁸⁶Sr strontium values) or to regions outside of the basins, which have higher ⁸⁷Sr/⁸⁶Sr values.

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287 2.3. The sites of Pech de l'Azé IV and Roc de Marsal

288 Pech de l'Azé IV (Pech IV) is a collapsed cave and part of a larger group of sites in a dry 289 tributary valley of the Dordogne River. The site was discovered and tested in 1952 by François 290 Bordes, and then trenched by an amateur archaeologist between 1953-1956. Bordes directed 291 excavations there between 1970 and 1977, exposing and area of 52 m² and an archaeological 292 sequence with a maximum depth of 4.5 m. For a history of research at Pech IV and neighboring 293 sites see Sandgathe et al. (2018). The lithic artefacts recovered played an important role in 294 Bordes' definition of several of the Mousterian *facies* (Bordes, 1975). Bordes excavated in 1 m² 295 squares following natural stratigraphy and archaeological features, and piece-plotting finds with 296 tape measurements. More recently, Dibble and McPherron's 2000-2003 excavations 297 demonstrated good stratigraphic agreement with Bordes' (Dibble et al. 2018). The specimens we 298 analyzed come from the 1970-1977 campaigns, specifically from Levels I2, Y and Z (the latter 299 being the two lowermost levels in the sequence). All these levels were assigned to the Typical 300 Mousterian (Bordes, 1975). Bordes' Level I2 corresponds to Dibble and McPherron's Layer 4C, 301 and Levels Y and Z collectively correspond to Layer 8 and are therefore lumped together here as

302 Level YZ. Layer 4C is characterized by abundant roof spall and occurrence of gelifraction, while 303 Layer 8 is characterized by abundant traces of burning, particularly burned bone, burned lithics, 304 ash, char (gelified organic matter) and reddened sediment (Goldberg et al., 2018). Bordes' 305 (1975) descriptions of Levels I2 and YZ are consistent with these observations. Two 306 thermoluminescence (TL) dates (68.5 ± 6.6 and 71.8 ± 6.7 thousand years [ka] before present 307 [BP]) place Layer 4C in the cold marine isotope stage (MIS) 4 (Richter et al., 2013), consistent 308 with the frost-related features (roof spall and gelifraction) mentioned above. This age is 309 consistent with the weighted mean OSL age of 68 ± 4 (Jacobs et al. 2016; Jankowski, 2018). 310 Layer 8 has a weighted mean age of 94 ± 4 based on OSL (Jacobs et al. 2016; Jankowski, 2018), 311 which is consistent with a TL age of 96±5 ka BP (Richter et al., 2013). Charcoal recovered from 312 Layer 8 comes from temperate woodland taxa including oak, birch, elm and hornbeam (Dibble et 313 al., 2009). Therefore, we assign the specimens from Level YZ to the interglacial MIS 5 (late MIS 314 5c and/or early MIS 5b). The fauna is consistent with these age determinations, being dominated 315 by reindeer in Layer 4C/I2 and by red deer in Layer 8/YZ (Hodgkins et al., 2016; Niven and 316 Martin, 2018).

317 Roc de Marsal (RDM) is located in the valley of a tributary of the Vézère River near the 318 town of Le Bugue. The site was first excavated by Jean Lafille between 1953-1971 (Lafille, 319 1961; Turq et al., 2017), leading to the discovery of a Neandertal child (Bordes and Lafille, 1962; Gómez-Olivencia and García-Martínez, 2019; Legoux, 1966). New excavations between 320 321 2004-2010 recognized 13 layers, with layers 10-2 containing significant Middle Paleolithic 322 remains (Reeves et al., 2019; Turq et al. 2008). Our samples come from layers 4, 7, 8 and 9 of 323 the new excavations. For chronological purposes, the sequence is divided into an upper part 324 dominated by Mousterian Quina technology (layers 4-2) and a lower part dominated by Levallois 325 technology with the base being richer in denticulates (layers 9-5). The two aggregates have been 326 dated through a combination of TL, OSL and infrared stimulated luminescence to ~ 49 ka BP 327 and ~70-65 ka BP, respectively (Guérin et al., 2017). Therefore, based on the luminescence 328 dating, the upper aggregate, including our samples from Layer 4, are associated with MIS 3, 329 while the lower aggregate, including our samples from layers 7-9, are associated with MIS 4. 330 However, the large mammal remains and phytoliths suggest that the upper aggregate formed 331 during colder conditions and the lower during milder and more humid ones (Hodgkins et al., 332 2016; Wroth et al., 2019).

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334 3. Materials and Methods

335 Samples for this study include archaeological specimens of Rangifer, Bison, Equus, Cervus, 336 *Capreolus* and rodent (*Microtus arvalis*) teeth, which were collected from the sites of Pech IV and 337 RDM. In addition to these samples, wine, soil, water and snail shells (Helix aspersa) were also 338 collected following standard protocols for developing baseline datasets for strontium (see Bentley, 339 2006; Evans and Tatham, 2004) throughout France to identify distinct isotopic regions and to 340 establish where herbivores may have been living and/or moving during tooth development (see 341 Table 3 and Figure 1; GPS coordinates available in Supplementary Online Material). One sample 342 location was the town of Les Eyzies, located 10 km from RDM and 30 km from Pech IV, which was used to help establish the range of ⁸⁷Sr/⁸⁶Sr values found close to the sites (Figure 1 Map A). 343 344 Modern snail shell samples were obtained from Mansle, located 158 km away from the 345 archaeological sites in the northern portion of the Aquitaine Basin (Figure 1 Map A). These 346 samples will be referred to throughout the paper as "local" values. Additional modern samples 347 were taken in the Alps, Pyrénées, and the edge of the Massif Armorican. The values obtained from

all samples collected will be compared to values published in Willmes et al., (2018). The geology
of France is complex being characterized by basins (Aquitaine and Paris basins) composed of
young sediments and uplifted hills, plateaus, and mountainous areas composed of older granites.
The presence of these distinctive regions increases the likelihood that mammal movement patterns
can be tracked across the landscape.



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Figure 1: A) Topographic map with the major geographic features of France discussed in the text with archaeological sites (black dots shown on all maps) and isotopic sampling sites collect for this paper indicated (red stars shown on all maps). B) Isoscape of bioavailable strontium in France (Willmes et al. 2014; data from IRHUM). C) Zoomed-in isoscape of bioavailable strontium in France focused on the Aquitaine basin. (Willmes et al. 2014; data from IRHUM). D) Isotope groups as determined by Willmes et al. (2018).

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362	Isotope sampling took place in the Archaeological Chemistry Laboratory (ACL) at Arizona
363	State University (ASU) and in the Facility for Isotope Research and Student Training (FIRST) at
364	Stony Brook University (SBU). A first set of samples (hereafter Set 1) including all modern Helix
365	shells $(n = 7)$, soils $(n = 5)$, water $(n = 8)$, wine $(n = 3)$, rodent teeth $(n = 3)$ and a portion of
366	archaeological herbivore teeth ($n_{samples} = 63$) was prepared for Sr isotope analysis in the Metals,
367	Environmental and Terrestrial Analytical Laboratory (METAL) at ASU and at the class 100 HEPA
368	filtered FIRST lab in the Department of Geosciences at SBU. A second set of samples (hereafter
369	Set 2) including archaeological herbivore teeth ($n_{samples} = 44$) was prepared in the Picotrace class
370	ten clean room at the Metal Geochemistry Center of the Department of Earth and Planetary
371	Sciences at Yale University.

The archaeological samples include molars and premolars of herbivores and rodents. Studies of dental development and eruption in ungulates demonstrate that these teeth develop over the first few years of life in these animals (see Table 1); thus, linearly-sampling enamel bands at intervals from each tooth has the potential to gain several 'snapshots' into range use during the period of tooth formation and to compare those data to the local bioavailable values where the teeth were found.

379 Table 1. Tooth crown formation and eruption of relevant teeth in large herbivores

			1	8	
	Taxon	Tooth	Enamel mineralization	Crown eruption	References
		M1	begins in utero	3-5 months	
	Rangifer	M2	starts <3.5 months	9-12 months	1,2,3,5,7,9
		M3	starts at 9-12 months	< 24 months	
-		M1	begins in utero	3-5 months	3,5
	Cervus	M2	starts <3 months	10-15 months	
		M3	starts at 9 months	15-29 months	
-	Equus	M1	birth to 23 months	8-12 months	5,6,7,8,10
		-			

	M2	7-37 months	20-26 months	
	M3	21-55 months	~3.5 years	
	P2	13-31 months	~2.5 years	
_	P4	19-51 months	~4 years	
	M1	begins in utero	9-12 months	
Bison	M2	birth to 13 months	18 months	4,7
_	M3	9-24 months	2.5-3 years	
Capreolus	P4	birth to 10 months	11-15 months	11

380

381 References for Table 1. ¹ Britton et al. (2011); ²Wu et al. 2012; ³Brown and Chapman, 1991;
⁴Gadbury et al., 2000; ⁵Hillson, 2005; ⁶Hoppe et al., 2004; ⁷Kelly, 2007; ⁸Levine, 1982; ⁹Spiess,
1979; ¹⁰Tomé and Vigne, 2003)

385

384

Enamel forms in two waves with different rates and geometry: secretion of a matrix (~20-30% of final mineral weight) and maturation (~70% of final mineral weight) (Green et al., 2017). During maturation, enamel is an open system and diffusion of elements occurs, resulting in averaging of the chemical signal (Balasse, 2003; Bentley, 2006; Fincham et al., 1999; Montgomery et al., 2010; Scharlotta et al., 2011). However, environmental chemistry is still reflected – albeit in attenuated form – throughout the tooth (e.g., Green et al., 2018; Uno et al., 2020).

392 Ideally, strontium isotope analyses would be conducted on individual mandibles or 393 maxillae with all molars present, or at least on second and third molar pairs (Britton et al., 2011); 394 however, it is difficult to find well preserved and associated dentition in archaeological 395 assemblages. At the time this study was undertaken, two partial mandibles were available from 396 RDM layer 4 for sampling, with mostly complete premolar and molar dentitions: specimens J17-397 2266 a bison, and J18-1379 a Rangifer. In these specimens, isotope samples were taken from all 398 available molars. In addition to these specimens, a Cervus mandible (F17-2485) was recovered 399 from Pech IV layer YZ that contained the second and third molars, and these teeth were sampled. 400 In order to expand the dataset to help establish whether or not different species were likely to have

401 inhabited ranges that included regions outside of the Aquitaine Basin, it was necessary to also 402 analyze single unassociated teeth. Teeth were chosen that appeared to be well preserved, 403 displaying few cracks in the enamel, and with as little discoloration as possible. Teeth were also 404 chosen based on the extent of wear. Teeth categorized as having no wear to early-wear were 405 prioritized; however, these teeth were not common and nearly half of the specimens sampled were 406 categorized as mid-wear, which means that strontium accumulated in these teeth during the first 407 few months of development could not be sampled due to wear.

At Pech IV a total of 11 individuals (12 teeth) were sampled. From layer I2 these specimens included: five *Rangifer* ($M^{1 \text{ or } 2}$, $M^{1 \text{ or } 2}$, M_1 , M_3 , P/M), two *Bison* (M^1 , M_2), and one *Equus* (M^1), which were excavated from units distributed throughout the cave (squares C13, D11, D14, D17, E11, F16). Specimens from layer YZ included: two *Cervus* (one partial mandible with M₂-M₃ and one isolated M^1) one *Equus* (M_2). These specimens were excavated from a unit located in the mouth of the cave (F16, F17, M15).

414 A total of 18 specimens (20 teeth) were sampled from RDM. From layer 4, these specimens 415 included one bison (partial mandible with M_1 and M_2), five *Rangifer* (one partial mandible with M₁ to M₃ row, and isolated M^{1 or 2}, M¹, M^{1 or 2}, M³), one Cervus (M¹) and two Equus (M² and P²) 416 417 excavated from just inside the cave mouth (squares J16, J17, J18, K16, K17). A single rodent tooth 418 from layer 7 was analyzed. Specimens analyzed from layer 8 included: two Rangifer (M₃, M₁), two Cervus (M1, M3), two Equus (P4, P/M), one Capreolus (P4) and one rodent excavated from 419 420 squares located just inside the cave mouth and on the terrace just outside the cave (G18, K17, K18,). For layer 9 sampled specimens included two Cervus (M³, M₁) and one rodent, recovered 421 422 from just inside and just outside the cave mouth (squares F18, G18, K17).

423 To determine if the teeth analyzed in this study have been diagenetically contaminated over 424 time, major, minor, and trace element analysis (referred to throughout as elemental analysis) was 425 run on a sub-set of herbivore teeth from each site. All archaeological materials are subject to 426 diagenetic processes, but these processes are not uniform over space or time, and therefore 427 specimens buried in different locations within the same site may undergo different chemical and 428 physical changes. A minimum of one specimen from each of the main archaeological layers 429 analyzed in this study (Pech IV layer I2 and YZ, RDM layer 4 and 8) were selected for elemental 430 analysis. An attempt was also made to select specimens from multiple excavation units within each 431 layer. The selected individuals therefore serve as proxies for contamination of other specimens 432 excavated from those same layers and units. In addition, from RDM one specimen from layer 4 433 and one specimen from layer 8 were selected to have both the enamel and the dentin analyzed. 434 This analysis was done to determine the level of contamination that occurred in the more porous 435 dentin portions of specimens as compared to the level of contamination that may have occurred in 436 the enamel.

437

438 3.1. Elemental analysis

Enamel or dentin powder was removed from specimens (see section 3.2; n = 12.3 mg each) at METAL, placed in a 15 mL centrifuge tube, and 960 μ L of twice distilled 5M nitric acid was then added to each tube along with 14 mL of Millipore water. All samples were analyzed using a Thermo Electron X-Series (with collision cell technology) Quadrupole Inductively Coupled Plasma Mass Spectrometer (Q-ICP-MS). Calibration standards was a combination of single-element standards designed to be similar in element ratios to bone and enamel. An internal element standard of Sc, Ge, In, and Bi was mixed online with all samples to correct for any change to instrumental sensitivity. A series of solutions with known concentrations (an internal
standard CUE-0001, a llama bone, and an external standard NIST-1400, an ashed bovine bone),
all of which were matrix-matched for archaeological bone, were measured repeatedly at regular
intervals throughout the analysis to monitor the accuracy and precision of the measurements.

450

451 3.2. Sample preparation and purification

452 All teeth (including rodent teeth) and modern snail shells were mechanically cleaned using a 453 Dremel Minimite-750 cordless drill with a tungsten carbide bur (Dremel engraving cutters 454 number 106). These samples were then rinsed with high purity deionized water (Millipore 18.2 455 megaohms) and submerged in an ultrasonic bath of Millipore water for 30 minutes. Due to the 456 potential for diagenetic contamination, the Dremel drill and cleaned carbide burs were used to 457 remove the outer enamel surface of each tooth (Budd et al., 2000; Hoppe et al., 2003). For snail 458 shell samples, each shell was drilled in half and both the outer and inner surface of the shell were 459 removed. Shells and rodent teeth were then placed in 10 ml of 0.8 M acetic acid for five to 460 twenty minutes and then rinsed with Millipore water. Once the teeth of large herbivores were 461 mechanically cleaned, sampling began by drilling horizontal lines into each tooth to collect 4-6 462 mg of enamel powder. Bands were drilled sequentially down the tooth, with the first band drilled 463 close to the occlusal surface and the last band close to the neck of the tooth. After each band of 464 enamel was drilled, the tooth was again placed in an ultrasonic bath of Millipore water for 10 465 minutes so that all powder adhering to the tooth was removed before the next sample was taken. 466 The resulting enamel samples were then analyzed under a light microscope to ensure that no 467 dentin was included. Sampling snail shell and rodent teeth began by placing the mechanically 468 and chemically cleaned specimens into crucibles in a furnace at 800° C for 14 hours and then

469 collecting the ashed samples. Set 1 (processed at ASU), tooth and snail shell samples were 470 dissolved in acid washed 1.5 mL centrifuge tubes with 0.5 mL of twice distilled 5M nitric acid 471 (HNO₃). Set 2 (processed at Yale) tooth samples were dissolved in acid washed Teflon beakers 472 with 1 mL 6.2 M hydrochloric acid (HCl) on a hotplate at 100° C overnight. After dissolution, 50 473 μ L of 30% hydrogen peroxide (H₂O₂) were added before evaporation (see below). 474 Soil samples were processed by first grinding each sample with a mortar and pestle. Once 475 ground, 4.0 grams of each sample were ashed in a furnace at 800° C for 14 hours. After ashing, 476 100 mg of each sample were reground with a mortar and pestle and placed in 15 mL centrifuge 477 tubes and dissolved in 5.0 mL of twice distilled 5M HNO₃ and 1.0 mL of hydrofluoric acid (HF). 478 In experiments run by Frank et al. (2022) different methods of soil leaching returned similar ⁸⁷Sr/⁸⁶Sr values, and where consistent to ⁸⁷Sr/⁸⁶Sr values sampled from vegetation at the same 479 480 sampling site. 481 Water and wine samples were processed by pipetting 20 mL of each water sample and 2 mL 482 of each wine sample into acid washed 20 mL Teflon beakers. Once samples were placed in the 483 beakers, 500 μ L of twice distilled 5 M HNO₃ were added to each wine sample. 484 After the steps above, all samples were evaporated on a hotplate. Twice distilled 3M HNO₃ (250 µL for Set 1 samples, 500 µL for Set 2 samples) was then added to each sample. Acid 485 486 washed columns were loaded with Eichrom SrSpec ion-exchange resin (30-50 µL for Set 1 487 samples, $\sim 70 \,\mu\text{L}$ for Set 2 samples). Resin was washed multiple times with twice distilled 3M 488 HNO₃ followed by Millipore water. The resin was then conditioned with twice distilled 3M 489 HNO₃. Next, the dissolved samples were loaded into columns, washed with twice distilled 3M 490 HNO₃ (500 µL for Set 1 samples, 1000 µL for Set 2 samples) and eluted with Millipore water 491 (500 µL for Set 1 samples, 1000 µL for Set 2 samples).

492

493 3.3. ⁸⁷Sr/⁸⁶Sr measurement

494 Set 1 samples were analyzed with a Thermo-Finnigan Neptune multi-collector inductively 495 coupled plasma mass spectrometer (MC-ICP-MS) at ASU's METAL. Solution concentrations 496 ranged between 3 and 158 ppb. Internal and external standards with varying concentrations were 497 run during analysis demonstrating that for samples with elemental Sr concentrations above 75 ppb the error of ⁸⁷Sr/⁸⁶Sr values was less than 0.00004. Sample solutions were diluted to achieve 498 499 a uniform concentration of 75 ppb Sr for measurement. 500 Set 2 samples were analyzed with a Thermo-Finnigan Neptune Plus MC-ICP-MS equipped 501 with an ESI Apex Omega sample introduction system at the Yale Metal Geochemistry Center. 502 Solution concentrations ranged from below reportable levels to 191 ppb Sr, with a single outlier 503 at 577 ppb Sr. All samples included in this analysis were measured at concentrations of 50 ppb. 504 For set 2, precision was evaluated through duplicate measurements of archaeological samples 505 $(\pm 0.00015, n=9)$ and accuracy through repeated measurements of an internal matrix-matched 506 standard (modern cow enamel, diluted from a concentrated stock solution and introduced at the same concentration as the samples) with known long-term ⁸⁷Sr/⁸⁶Sr (0.71643, average difference 507 508 $= -0.00010 \pm 0.00006$, n=4). 509 For both sets, corrections were applied to account for the isobaric interference of ⁸⁶Krypton (Kr) on ⁸⁶Sr (Kr being present as a trace impurity in the Argon gas used in the plasma) and ⁸⁷Rb 510

511 on ⁸⁷Sr (Rb being present in small amounts after imperfect column purification). The abundance

512 of ⁸⁶Kr was estimated from the measured abundance of ⁸³Kr, while the abundance of ⁸⁷Rb was

513 estimated from the measured abundance of ⁸⁵Rb accounting for instrumental mass fractionation

514 (Horsky et al., 2016; Konter and Storm, 2014). Measurements were corrected for drift using an

515 external intra-elemental correction (sample-standard bracketing) with NIST SRM 987 as the 516 bracketing standard.

517

518 4. Results

519 4.1. Elemental analysis

520 The average calcium-to-phosphorus ratios (Ca/P) measured on NIST-1400 over the course of the 521 analysis were 2.10 ± 0.06 (2σ , n= 9). This value is accurate and precise when compared to the 522 certified calcium and phosphorus values published by the National Institute of Standards and 523 Technology (Ca/P= 2.13 ± 0.03 , 2σ). 524

Results from elemental analysis demonstrate that most enamel samples were not

525 diagenetically contaminated, since Ca/P from enamel samples (mean Ca/P = 2.14 ± 0.08 , 2σ ,

526 n=12) closely match ratios found in modern teeth (~ 2.14) (Table 2) (Burton, 2008;

527 Papageorgopoulou and Xirotiris, 2009; Sillen, 1989). The one exception is a second molar from

an Equus specimen (F16-3786) from Pech IV, layer YZ, which has an elevated Ca/P (⁴⁴Ca/³¹P 528

529 =2.24), suggesting some contamination may have occurred in this specimen. As expected from

530 archaeological samples, Ca/P in dentin were elevated demonstrating that dentin is more

531 susceptible to contamination than enamel (Budd et al., 2000). Other tests for diagenetic

532 contamination included measuring concentrations of neodymium (Nd) and uranium (U). Results

533 reveal that none of the enamel samples have undergone contamination by uranium or

534 neodymium as the U/Ca and the Nd/Ca values are within one standard deviation of ratios

535 measured in the NIST 1400 standard. Additionally, U concentrations in all tooth samples were <

536 0.1 ppb, which is within the range that has been reported from modern bones (Burton, 2008;

537 Grimstead et al., 2018; Williams and Marlow, 1987). It has been noted that since U and Nd are

538 universally present in soils, finding low values of these elements in archaeological samples is a

539 good indication that the samples have largely been resistant to chemical alteration (Burton,

540 2008). Thus, these results suggest that strontium measured in archaeological teeth from RDM

and Pech IV (with the possible exception of specimen F16-3786) will likely reflect biogenic and

542 not diagenetic values.

543

544 545

546

Site, Layer	Specimen	Sample	Ca/P	U/Ca	Nd/Ca
	J18-4657 Equus	ACL-2328 enamel	2.09	1.73E-07	5.19E-07
	K17-1338 <i>Cervus</i>	ACL-2340 dentin	2.79	3.02E-06	3.50E-07
RDM, L4	K17-1338 <i>Cervus</i>	ACL-2340 enamel	2.15	2.55E-07	5.12E-07
	J16-1572 Rangifer	ACL-2337 enamel	2.10	1.46E-07	1.11E-06
	K16-2238 Equus	ACL-2341 enemel	2.15	1.64E-07	1.23E-06
	K17-2382 <i>Rangifer</i>	ACL-2330 dentin	2.35	4.18E-06	4.55E-06
	K17-2382 <i>Rangifer</i>	ACL-2330 enamel	2.11	6.07E-07	7.02E-07
RDM, L8	G18-1882 <i>Rangifer</i>	ACL-2329 enamel	2.17	1.20E-07	3.46E-06
	K17-1418 <i>Equus</i>	ACL-2334 enamel	2.10	1.03E-07	3.03E-07
RDM, L9	G18-6729 <i>Cervus</i>	ACL-3434 enamel	2.13	3.87E-07	3.63E-07

Deck I 12	E11-3125 Rangifer	ACL-2325 enamel	2.14	8.13E-06	2.85E-06
Pecn L 12	D11-3951 <i>Rangifer</i>	ACL-2338 enamel	2.15	3.05E-07	9.61E-07
Deck I V7	F17-2845 <i>Cervus</i>	ACL-3435 enamel	2.10	1.90E-07	4.58E-07
	F16-3786 <i>Equus</i>	ACL-3436 enamel	2.24	9.60E-07	2.07E-07
CUE-0001 average		n= 5 (2σ)	$2.13\pm\!\!0.06$	2.35E-07± 1.96E-07	2.91E-08± 1.85E- 07
NIST 1400 average		n=9 (2σ)	2.10 ±0.06	1.03E-07 ± 1.17E-05	5.09E-07 ± 1.98E-06

548

549 Table 2: Results from major, minor, and trace element analysis

550

551

552 4.2. Radiogenic strontium isotope analysis

553 For Set 1, data quality was evaluated through repeated measurements of NIST SRM 987, which

indicate good precision (87 Sr/ 86 Sr = 0.710265 ± 0.000014, 2 σ , n=18,) and agreement with

accepted (0.71025) and certified (0.71034 \pm 0.00026) values for the standard (Faure and

556 Mensing, 2005; Moore et al., 1982). For Set 2, precision was evaluated through duplicate

557 measurements of archaeological samples (±0.00015, n=9) and accuracy through repeated

558 measurements of an internal matrix-matched standard (modern cow enamel) with known long-

559 term 87 Sr/ 86 Sr (0.71643, average difference = -0.00010 ±0.00006, n=4).

Sample	⁸⁷ Sr/ ⁸⁶ Sr Range	Mean	σ	n Samples
Fossil Herbivore Samples				
All herbivores from Pech IV and RDM	0.7087-0.7138	0.7100±0.0009	1	107
Cervus	0.7097-0.7138	0.7111±0.0013	1	18
Capreolus	0.7096-0.7099	0.7098±0.0001	1	2
Bison	0.7096-0.7108	0.7101±0.0005	1	12
Equus	0.7087-0.7115	0.7197±0.0006	1	38
Rangifer	0.7091-0.7103	0.7097±0.0002	1	37
Modern Samples: soil, water,	, wine, <i>Helix shells</i>			
Local values for Les Eyzies and Mansle (includes <i>Microtus</i> teeth from archaeological layers)	0.7087-0.7092	0.7090±0.0002	1	7
Alps	0.7072-0.7250	0.7132±0.0067	1	8
Granitoid regions of the Alps: Chamonix Lacs des Gaillands, La Fontaine, Saint Amour	0.7113-0.7250	0.7186±0.0055	1	4
Massif Armorican	0.7137-0.7138	0.7137	1	2
Pyrénées	0.7084-0.7155	0.7108±0.0027	1	8

560

561 **Table 3: Summary of the** ⁸⁷**Sr**/⁸⁶**Sr obtained in this study**

562

563 Table 3 summarizes the ⁸⁷Sr/⁸⁶Sr values obtained from soil, water, wine, and snail shells for this 564 study. The ⁸⁷Sr/⁸⁶Sr values obtained from these modern environmental samples are consistent 565 with the ⁸⁷Sr/⁸⁶Sr isoscape published by Willmes et al. (2018), which uses lithology to maximize 566 the difference between geological units. This approach averages isotopic values to reduce noise, 567 while sacrificing specificity (see Figure 1, and Table 3). This is a conservative but useful 568 approach for animal movement studies because even non-migratory animals are likely to forage 569 over a variety of lithologies even with a small home range. The local values obtained from Les 570 Eyzies and Mansle (0.7087-0.7092) broadly fall within Willmes and colleagues' Isotope Group 571 2 (0.7072–0.7115), which includes carbonaceous rocks, including those characterizing the 572 Aquitaine and Paris basins; however, some of our local values are compatible with Isotope 573 Group 3 (0.7076–0.7170), including clastic sediments within these basins, although Group 3 574 includes higher values (the group also includes rock types found in the mountains). Values from 575 samples we collected in the Alps (0.7072-0.7250) span Isotope Groups 3 and 4 (0.7084–0.7252), 576 the latter including rock types found in the mountains, most notably granite. Values from the 577 samples we collected from the Massif Armorican (0.7137-0.7138) and the Pyrénées (0.7084-578 0.7155) span Isotope Groups 3, 4 and 5 (0.7155–0.7213), which collectively include the various 579 rock types found in these mountain ranges.

580 The range of values obtained on herbivore teeth sampled in this study (0.7087-0.7138)581 spans and exceeds our local values and those of Isotope Group 2 from Williams et al. (2018) to 582 include ratios characteristic of mountainous regions, indicating that at least some taxa or 583 individuals can be confidently labelled as non-local. Rangifer values obtained in this study 584 (0.7091-0.7103) overlap (Wilcoxon rank sum test W = 726, p = 0.51) with those obtained at 585 Jonzac (0.7095-0.7104, mean 0.7098 \pm 0.0002, 1σ , n = 36; Britton et al., 2011) and both fall 586 within Isotope Group 2 variability, while exceeding the variability of our local environmental 587 samples. Our Bison values (0.7096-0.7108) exceed local values as defined here but fall within 588 Isotope Group 2. These values are meaningfully different compared to Rangifer from the same 589 sites (W = 304, p = 0.058) as well as *Bison* from Jonzac (0.7090-0.7092, mean 0.7091±0.0001,

590	1σ , $n = 12$; Britton et al., 2011) (W = 144, $p < 0.00001$). Equus values (0.7087-0.7115) exceed
591	local values but fall within Isotope Group 2, and are distinct from those of <i>Bison</i> ($W = 137.5$, $p =$
592	0.04) but not <i>Rangifer</i> (W = 627.5, $p = 0.43$) at the same sites. <i>Cervus</i> yields the largest range of
593	variation in our sample (0.7097-0.7138, $\Delta = 0.0041$), exceeding the variability of Willmes
594	Isotope Group 2 (0.7072-0.7115) and indicating with confidence that this taxon was in areas with
595	older bedrock, which may include the Massif Central, Pyrénées, or perhaps areas in the
596	Armorican Massif located in northeastern France bordering the Bay of Biscay (Figure 1). Cervus
597	values are statistically different from those of Rangifer, Bison and Equus from the same sites (W
598	= 605.5, 173.5, 592, respectively, all $p < 0.01$). Finally, two values for <i>Capreolus</i> (0.7096 and
599	0.7099) exceed local variability but fall within Isotope Group 2. Despite the very small sample of
600	Capreolus, this (likely) non-migratory taxon appears indistinguishable from Rangifer and Bison
601	in our study.



Figure 2. ⁸⁷Sr/⁸⁶Sr measured on archaeological enamel from RDM and Pech IV, grouped
by age. The black dashed lines mark the boundaries of local baseline as measured in this
study. The red dashed lines mark the boundaries of Isotope Group 2 as defined by Willmes
et al., (2018).

607

608 Breaking down our results by MIS reveals temporal trends in the data (Figure 2). When 609 all taxa are considered together (rodents excluded), there are statistically significant differences 610 across MIS groups (Kruskal-Wallis, $\chi^2 = 22.722$, p < 0.0001). These differences are driven by 611 MIS 5 at Pech IV (which has the smallest sample size = 9) being different from all other groups 612 (Pairwise W, p < 0.001 [Benjamini and Hochberg FDR correction for multiple testing here and 613 below]). The lack of Rangifer samples from the MIS 5 at Pech IV (i.e., layer YZ) is itself 614 meaningful and it is not an artifact of biased sample selection, since the zooarchaeological 615 analysis of layer YZ yielded a single *Rangifer* specimen but a rich assemblage of *Cervus* 616 (Hodgkins et al., 2016). 617 Differences in Rangifer values across MIS 4 at Pech IV and RDM and MIS 3 at RDM approach, but do not reach, statistical significance (Kruskal-Wallis, $\chi^2 = 5.4872$, p = 0.06). Bison 618 values differ between MIS 4 at Pech IV and MIS 3 at RDM (W = 0, p = 0.005), but in both cases 619 620 remain well within the variability of Isotope Group 2. MIS 5 Equus at Pech IV has clearly high, non-local values, but the other groups show differences as well (Kruskal-Wallis, $\chi^2 = 8.0298$, p =621 622 0.02), driven by MIS 4 *Equus* at RDM being different from MIS 3 *Equus* at the same site 623 (Pairwise W, p = 0.02). MIS 3 Cervus at RDM has clearly low, more local values than older 624 individuals, while MIS 4 and 5 Cervus at RDM and Pech IV, respectively, are indistinguishable

625 (W = 16, p = 0.22).

626 Measurements taken at different intervals along the crown heights of each tooth have 627 been plotted in Figures 3-6 for reindeer, bison, horse and red deer respectively. *Rangifer* display 628 some intra-tooth variability (i.e., trends from lower to higher, or higher to lower, values between 629 the two or three points measured in each of the teeth), although all measured values align with 630 bioavailable values determined for the Aquitaine and Paris Basins. While a higher resolution 631 sampling approach including multiple teeth from the same individual would be required to better 632 understand the seasonal movements these animals may have been undertaking, these data from 633 two or three distinct portions of the crown suggest that reindeer range may have been restricted 634 to the basins at this time. This is also the case with the *Bison*, where some intra-tooth variability 635 is displayed, but all values are within the range of values found in the basins. A single horse specimen, and several red deer display some more elevated ⁸⁷Sr/⁸⁶Sr values, with a single *Cervus* 636 637 M3 (RDM Layer 4, J18-3156) displaying a directionality from high values near the neck 638 (youngest enamel) to lower values near the occlusal surface (oldest enamel) (Figure 6). This 639 pattern shows that this individual moved from a geologically younger substrate to a geologically 640 older substrate. Assuming a late May-early June birth and start of M3 mineralization around nine 641 months of age (Brown and Chapman, 1991), this pattern could be consistent with a springtime movement from winter pastures at lower elevations (lower ⁸⁷Sr/⁸⁶Sr) to summer pastures deeper 642 into the mountains (higher ⁸⁷Sr/⁸⁶Sr). Higher resolution sampling (i.e. of multiple teeth, with a 643 644 great number of samples taken at closer intervals) would add clarity to this pattern and 645 determined if these movement patterns were seasonally repeating, as has been highlighted for 646 *Cervus* in the south of France during MIS 6 (Barakat et al. 2023).

647





661 on a single specimen. Square symbols denote an individual (RDM Layer 4, J17-2266) from

- 662 which multiple molars where sampled. The *Bison* samples all fall outside the range of
- 663 ⁸⁷Sr/⁸⁶Sr values for the local sample (0.7087-0.7092) but within Isotope Group 2 (0.7072-
- 664 **0.7115) from Willmes et al. (2018).**









- 675 measurements taken on a single specimen and square symbols denote an individual (F17-
- **2845) for which both M2 and M3 were sampled. Orange and blue shading denote estimated**
- 677 <u>summer and winter seasons respectively, assuming a May birth (Landete-Castillejos et al.,</u>
- **2005), tooth formation times after Brown and Chapman (1991), and unworn crown heights**
- 679 of ~27, ~30 and ~31 mm for M1, M2 and M3, respectively (Steele and Weaver, 2012). All of
- 680 the *C. elaphus* samples fall outside the range of ⁸⁷Sr/⁸⁶Sr values for the locally collected
- 681 sample (0.7087-0.7092) and three samples fall above the Isotope Group 2 range (0.7072-
- **0.7115) of Willmes et al. (2018).**

- **5. Discussion**

687	The comparison of ⁸⁷ Sr/ ⁸⁶ Sr from archaeological herbivore teeth and the bioavailable
688	⁸⁷ Sr/ ⁸⁶ Sr across France enables a more detailed analysis of the movement of these animals on the
689	landscape. The Rangifer sampled here appear to have been moving within the Aquitaine, and
690	perhaps Paris, basins of France. The same applies to Bison and the single Capreolus studied here,
691	which had the smallest ⁸⁷ Sr/ ⁸⁶ Sr range. Interestingly, non-migratory behavior in <i>Bison</i> is
692	documented by 87 Sr/ 86 Sr also for a relatively large (n = 25) group of individuals from eastern
693	Ukraine, dated to 18.5 ka BP (Julien et al., 2012). Equus was also mostly available within the
694	French basins, although a single individual (specimen F16-3786 from Pech IV, layer YZ dating
695	to MIS 5) displays more radiogenic values consistent with ranging in an area between a basin
696	and the mountains, though as noted in the results this individual may have been diagenetically
697	altered because it has an elevated Ca/P (44 Ca/ 31 P =2.24), so this result should be viewed with
698	caution. Finally, <i>Cervus</i> is the only taxon with ⁸⁷ Sr/ ⁸⁶ Sr values outside the range observed
699	in Rangifer, Bison, Equus, and Capreolus, thus they seem to have had a different foraging
700	pattern than the other species. Higher mobility in Cervus compared to sympatric Equus
701	hydruntinus was inferred from ⁸⁷ Sr/ ⁸⁶ Sr at Terminal Pleistocene sites in central Italy, although in
702	that case neither taxon appeared to forage in the mountains (Pellegrini et al., 2008). However, as
703	indicated by the laser ablation data recently obtained for Cervus from the site of Lazaret, strongly
704	seasonal strontium signals can be the result of short-distance altitudinal migrations, rather than
705	necessarily long-distance migrations (Barakat et al. 2023). It is also possible that Cervus was
706	foraging closer to the coast by the Armorican Massif bordering the Bay of Biscay.
707	While these results do not speak to the distance over which these species were moving
708	(they could have been moving from the Aquitaine basin to the Paris basin), the fact that most do

not seem to have spent time in the mountains is at odds with many hypotheses of prey movement

710 patterns in southwestern France. Today, regional movements like this observed in *Rangifer* and 711 *Bison* have been tied to mosaic habitats of pine and spruce woodlands and boreal forests. Thus, 712 these results can be used to assist paleoenvironmental reconstructions of southwestern France. 713 When the radiogenic strontium isotope values measured in *Rangifer* teeth at RDM, Pech 714 IV and from Britton et al.'s (2011) study at Jonzac are compared, the range of values found are 715 nearly identical at all three sites. These results suggest that *Rangifer* where foraging over similar 716 geological units over time and climatic cycles. The intra-tooth sampling conducted by Britton et 717 al. (2011) found that those movements were likely to have been seasonal. The data provided here 718 could be complementary to that study, perhaps reindeer were undertaking seasonal movements 719 between different regions within the basins of France. Further work is needed to characterize the 720 seasonal nature of *Rangifer* behavior in this part of France during the Late Pleistocene and to explore whether this extended into the Paris Basin. Given the common ⁸⁷Sr/⁸⁶Sr values found in 721 722 both basins, complementary methods such as sulfur or lead isotope analyses may be needed. 723 Britton and coworkers (2011) also analyzed radiogenic strontium isotope ratios from the 724 second and third molars of a bison recovered from Jonzac. Results from this analysis showed that 725 there was little variation across the tooth, suggesting that *Bison* remained within one geological 726 region during tooth development, and unlike the reindeer, did not show evidence of seasonal 727 range use. The mobility pattern inferred from the Bison at Jonzac is consistent with the findings 728 of the current study, although one Bison from Pech IV does appear to have foraged in a more 729 radiogenic region than other *Bison* at the site. Analyses of *Bison* mobility from all three sites 730 have interesting implications for the Neandertals who would have hunted them. If Bison truly 731 were living within more restricted regions, they would have been present for Neandertals to hunt 732 year-round. Bison are large herbivores and a valuable resource and may have limited the need

for Neandertals to undertake long distance movements. An analysis of dental eruption and wear,
as well as aging of fetal animals, throughout the sequence at Pech IV suggests in fact that the site
was occupied during multiple seasons (Niven and Martin, 2018).

Finally, our data suggests that animals hunted during MIS 5 were either procured farther away from Pech IV and RDM compared to MIS 4 and 3 fauna, or at least lived part of their lives farther from these sites.

739

740 **6. Conclusion**

741 An analysis of variation in the radiogenic Sr isotope ratios of Bison, Equus, Rangifer, and 742 Capreolus suggest that all of these animals roamed within the basins of France and would have 743 been available for Neandertals to hunt which is important knowledge for reconstructions 744 Neandertal mobility and strategies. Of course, factors other than availability, could have 745 influenced early hominin hunting behaviors, including the seasonal conditions of animals (both 746 their body mass/fat content) and the condition of their pelts (Britton et al. 2023). For instance, at 747 Abri du Maras, a lack of seasonal movement in *Rangifer* in level 4.1 was combined with a 748 strongly seasonal exploitation pattern – indicating that Neandertals were only either using the 749 site seasonally or that their preference for reindeer at that time was strongly seasonal due to the 750 aforementioned factors (Britton et al. 2023). Cervus on the other hand may have ventured further 751 afield, towards the Atlantic coastline or into the mountains, which could indicate that 752 Neandertals ventured into these different ecotones to hunt them during MIS 4 and 5. 753 We hope that this study is useful to future researchers who can utilize methods beyond

those applied in this paper to pin down the distances each of these species traveled and/or their seasonal range use within the basins. Higher resolution sampling is necessary to see if subtle

756 repeating isotopic signals exist that could indicate movements within the basins due to their 757 relatively homogenous isotopic profile. In addition, and due to this homogeneity, complementary 758 methods, such as sulfur isotope analysis may still be required to identify range use that spanned 759 across the two basins and differences in total home range size between different species and 760 individuals (e.g., Britton et al. 2023). Given the anomalous pattern of Cervus reported in this 761 study, we hope future studies will characterize where summer and winter ranges of Cervus may 762 have been, and to determine whether or not these overlapped with the sites during seasonality of 763 site use/procurement in order to assess whether *Cervus* was a seasonally available resource 764 locally or whether Neandertals extended their hunting ranges to exploit this species. 765 The results of this analysis add important information about biogeography of key 766 Neandertal subsistence species and hint at potential changes in their range use over time, 767 contributing important foundational data to future studies and bringing additional understanding 768 to the ecological context of Neandertal subsistence practices.

769

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Figure 4

0 20 P4 40 60 0 80 20 40 20 0 80 60 40 Distance from occlusal (mm) P2 0 80 M2 60 0 80 20 M 40 60 80 ⁸⁷Sr/⁸⁶Sr 0.7104 0.7102 0.7114 0.7112 0.7110 0.7100 0.7092 7607.0 0.7094

Click here to access/download;Figure;Figure 5 HorseToothProfiles2 300 dpi.jpeg ≛

Figure 5



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Declaration of competing interest

The authors declare that they have no known competing financial interests or personal relationships that could have appeared to influence the work reported in this paper. The authors did not use any AI programs to write this paper.

Author Contributions

Each of the authors in the author list have contributed in the following ways:

Jamie Hodgkins: Conceptualization, Data curation, Formal Analysis, Funding acquisition, Investigation, Methodology, Project administration, Writing – original draft, Writing – review & editing

Alex Bertacchi: Formal Analysis, Methodology, Visualization, Writing - review & editing

Kelly J. Knudson: Conceptualization, Methodology, Writing - review & editing

Troy Rasbury: Conceptualization, Methodology, Writing - review & editing

Julia I. Giblin: Methodology, Resources

Gwyneth Gordon: Conceptualization, Methodology, Writing - review & editing

Ariel Anbar: Resources

Alain Turq: Resources

Dennis Sandgathe: Resources Writing – review & editing

Hannah M. Keller: Visualization, Writing - review & editing

Kate Britton: Conceptualization, Methodology, Writing - review & editing

Shannon P. McPherron Conceptualization, Methodology, Writing - review & editing

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